

# Challenges in Evaluating the Influence of the Ocean on Atmospheric Composition



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*GasEx-98, North Atlantic Ocean*

# Outline



*RB-99 (Phase), North Pacific Ocean*

- Fundamentals
- Tools and Toys
- Early Examples
  - $\text{N}_2\text{O}$
  - CFCs
  - $\text{CH}_3\text{CCl}_3$
- $\text{CH}_3\text{Br}$
- $\text{CCl}_4$
- Possibilities . . .

# Behavior of a Gas in the Atmosphere

$$\frac{dG}{dt} = \textit{Emission Rate} - \textit{Loss Rate}$$

$$\frac{dG}{dt} = E - kG$$

$$k = \frac{1}{\tau} = \frac{1}{\tau_{strat}} + \frac{1}{\tau_{trop}} + \frac{1}{\tau_{ocean}} + \frac{1}{\tau_{soils}} + \dots ,$$

# Air sea exchange formula

Kinetics

$$Flux = \frac{K_w A}{H} (p_w - p_a)$$

Thermodynamics

Ultimate Drivers: Partial pressure difference  
Wind speed  
temperature  
salinity



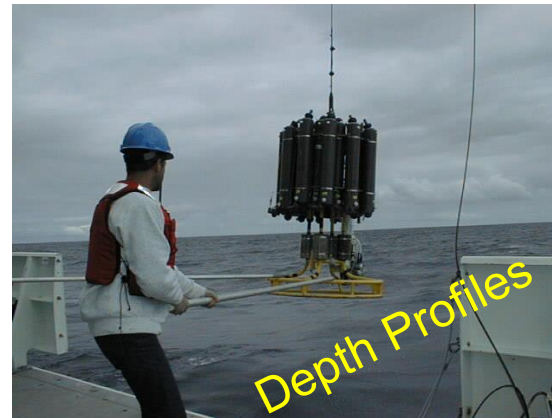
# Tools



Air Sampling



Equilibrator



Depth Profiles



Checking



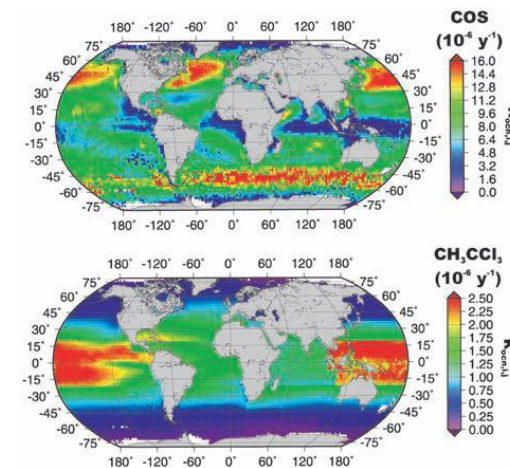
And checking



Controlling Delivery

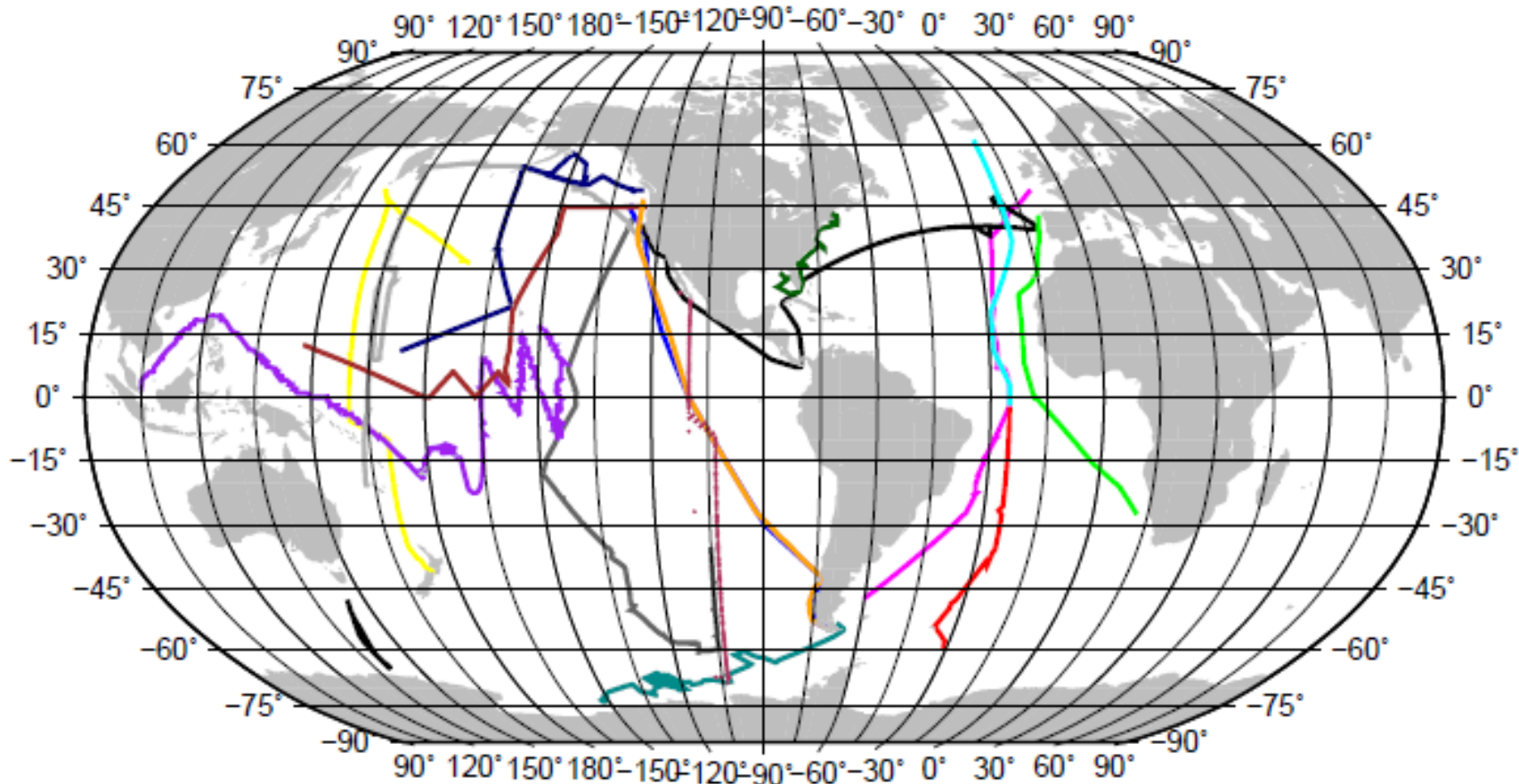


Analysis



$$\frac{dM_a}{dt} = S_A + F_I - \frac{K_w A}{H} (p_w - p_a) - k_{OH} p_a M - k_s p_a M$$

# Research Cruises 1987-2010



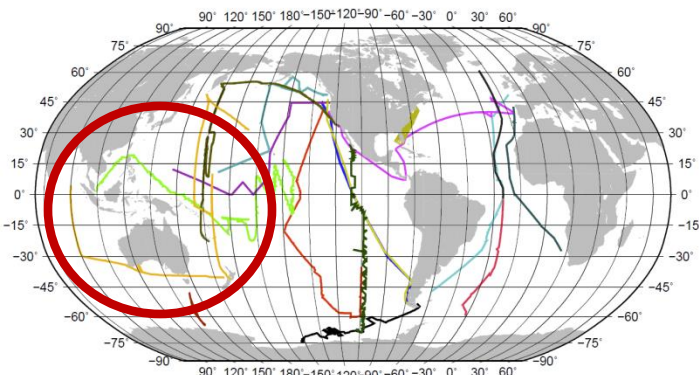
SAGA2 (1987), RITS89 (1989; also P18), SAGA3 (1990), OAXTC (1990; also P13),  
BLAST1 (1994), BLAST2 (1994), BLAST3 (1996), GasEx98 (1998), RB9906 (1999),  
CLIVAR01 SR3 (2001), A16N (2003), A16S (2005), PHASE (2004), P18 (2008),  
HalocAST A (2009), GOMECC (2010), HalocAST-P (2010)



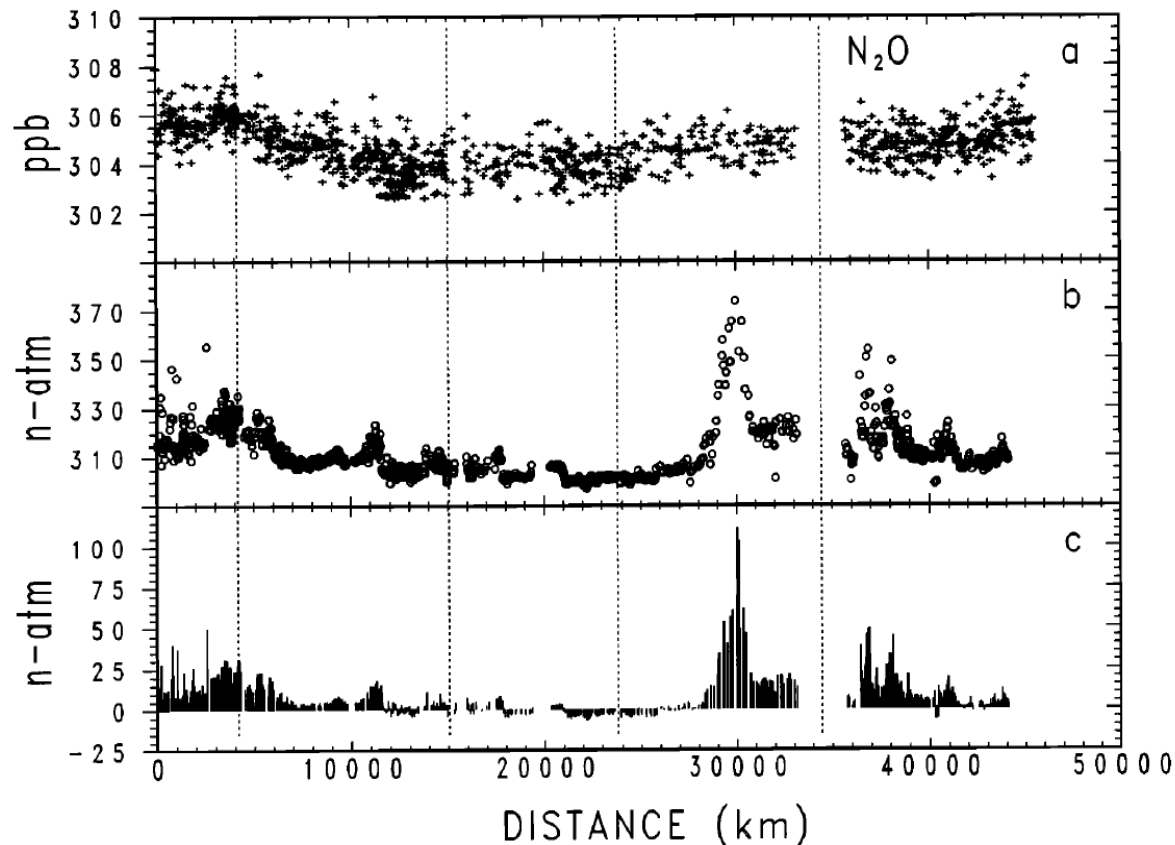
# N<sub>2</sub>O Emissions (1987)



Akademik Korolev



- West Pacific and East Indian Oceans
- Emissions associated with upwelling
- N<sub>2</sub>O produced at mid-depth

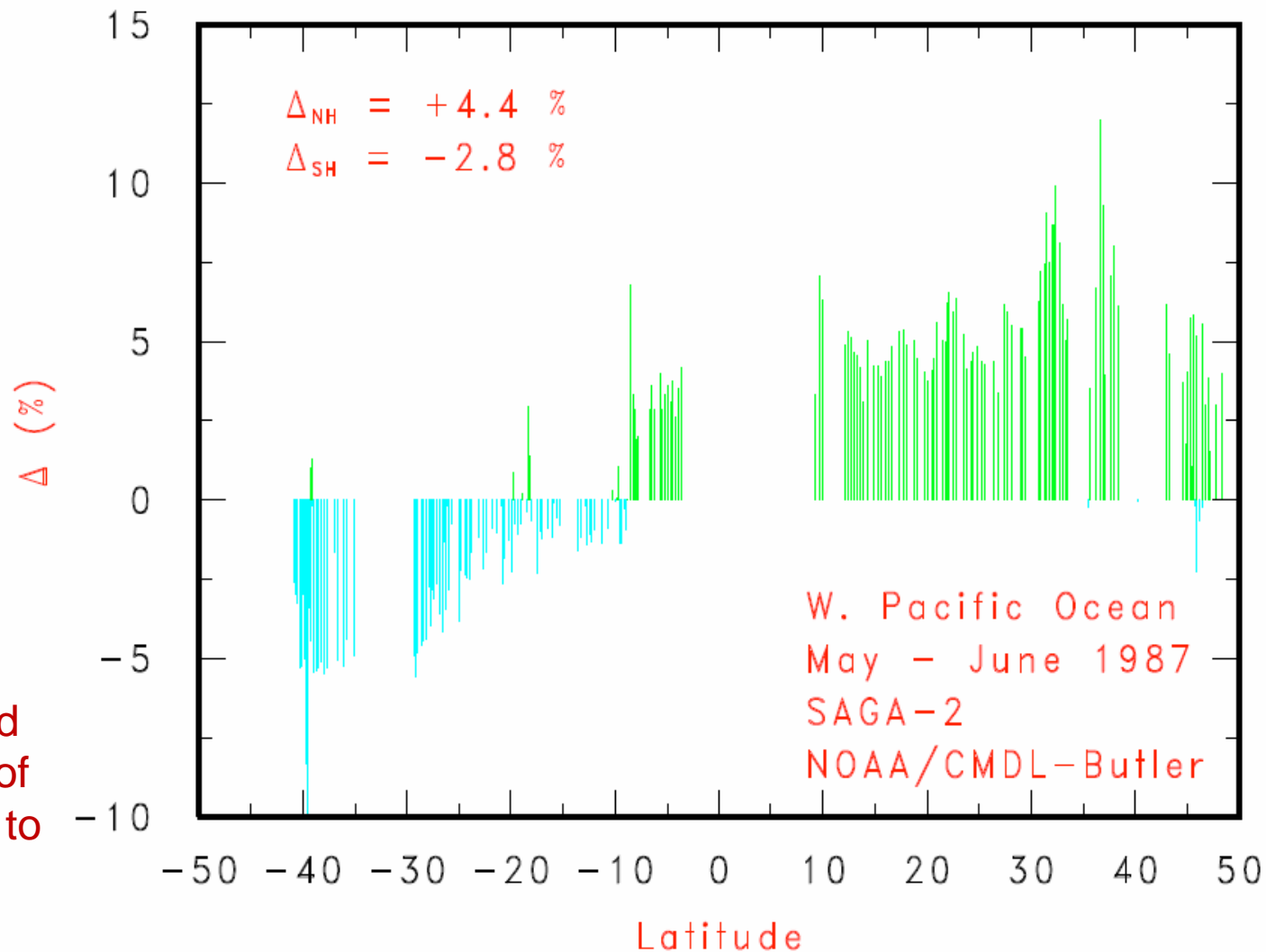


10-20% of atmospheric N<sub>2</sub>O  
comes from the ocean

# Physical Effects – “Inert” CFC-11 (1987)

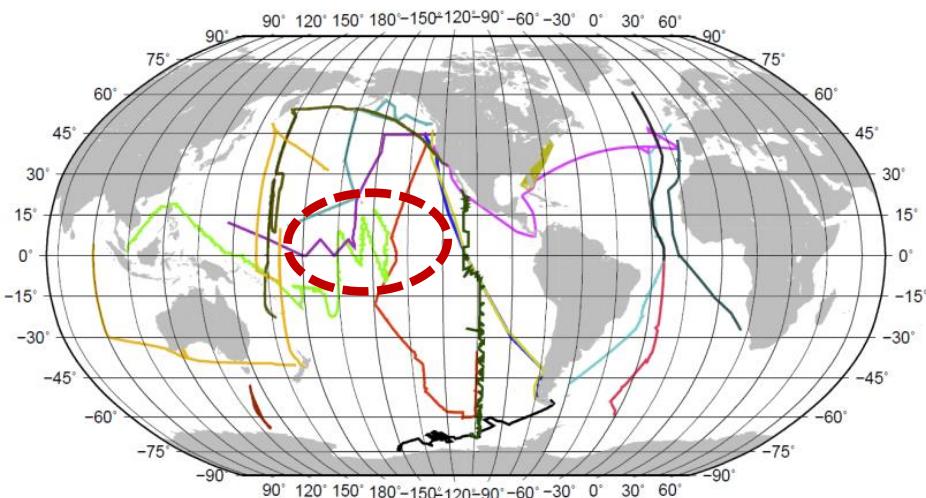
- Warming & Cooling
- Transport
- Mixing
- Air Injection

Much of an observed saturation anomaly of any gas can be due to physical effects.

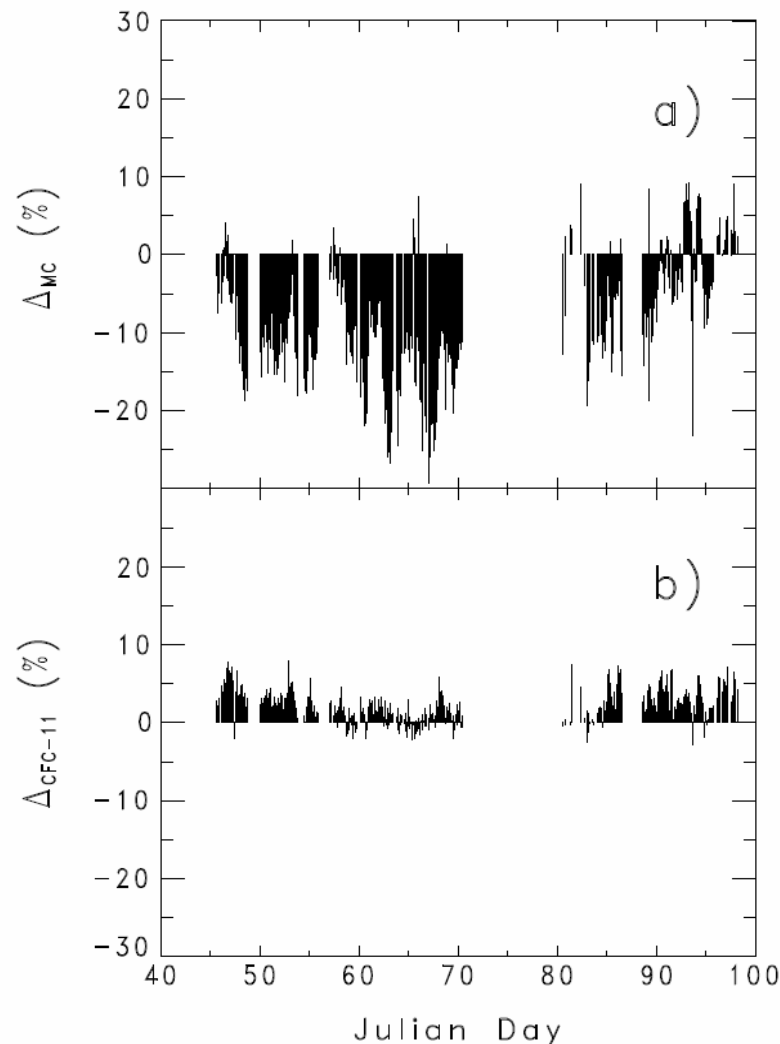




# Chemical Consumption – $\text{CH}_3\text{CCl}_3$ (1990)



- 5-11% of atmospheric  $\text{CH}_3\text{CCl}_3$  is consumed by hydrolysis in surface seawater
- This finding reduced the atmospheric lifetime of  $\text{CH}_3\text{CCl}_3$
- And, it increased the estimated lifetime of all other gases consumed by tropospheric OH
- And it changed its ozone-depletion potential



And then came  $\text{CH}_3\text{Br}$  . . .

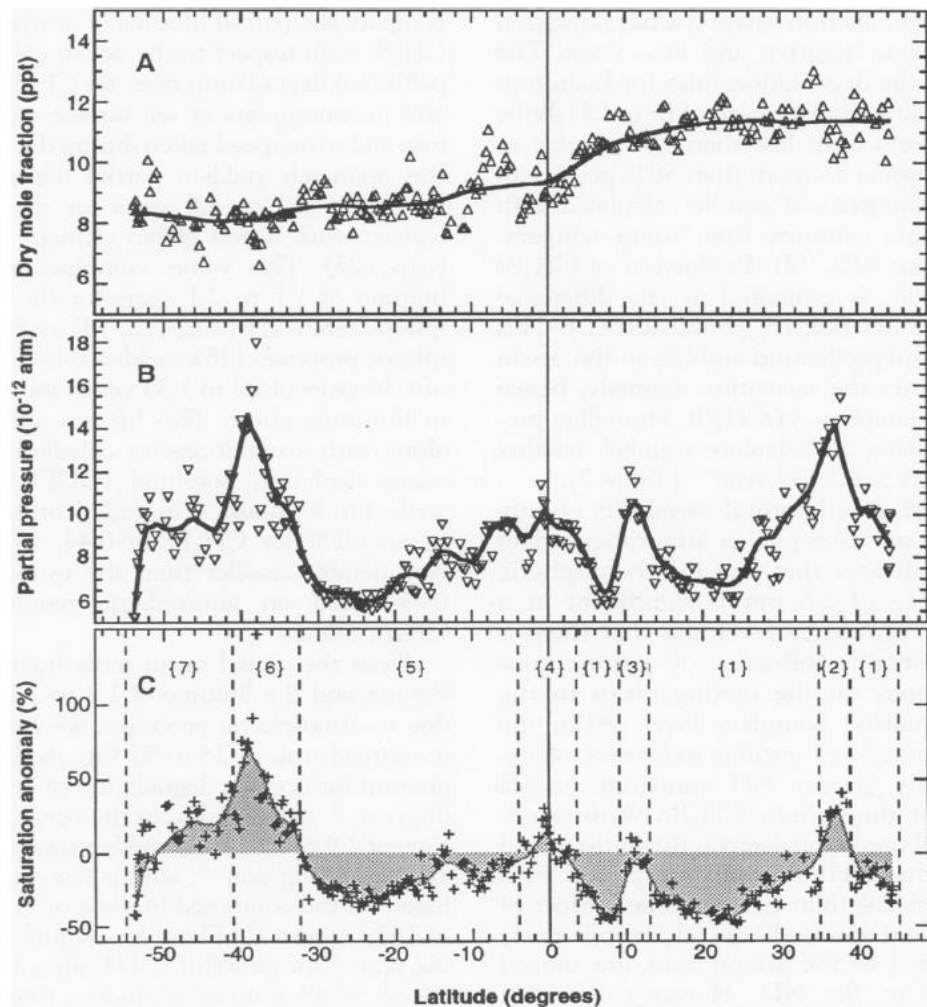
Where we really learned  
something . . .

# Properties of and issues with $\text{CH}_3\text{Br}$

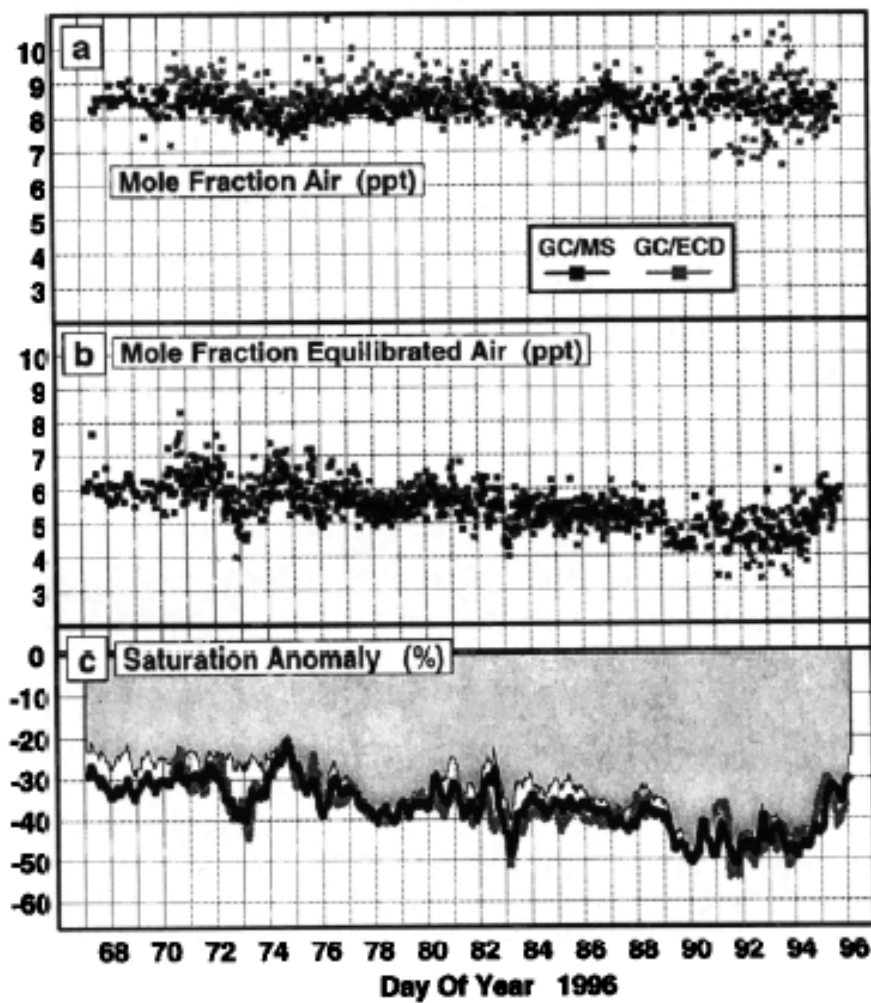


- Consumed in the surface ocean by nucleophilic displacement of Br by  $\text{Cl}^-$  (10-15%/d)
- Produced in surface seawater by microorganisms at similar rates (**Ocean once believed to be hugely supersaturated.**)
- Consumed in surface seawater by microorganisms at similar rates
- Surface fluxes reflect the sum of all of these
- How to separate them and get reliable estimates?
- Other questions:
  - Early issues with sampling and analysis reliability
  - Human emissions are significant for ozone depletion, but...
  - Only about 1/3 of the  $\text{CH}_3\text{Br}$  in the atmosphere is (was) anthropogenic

# Observations



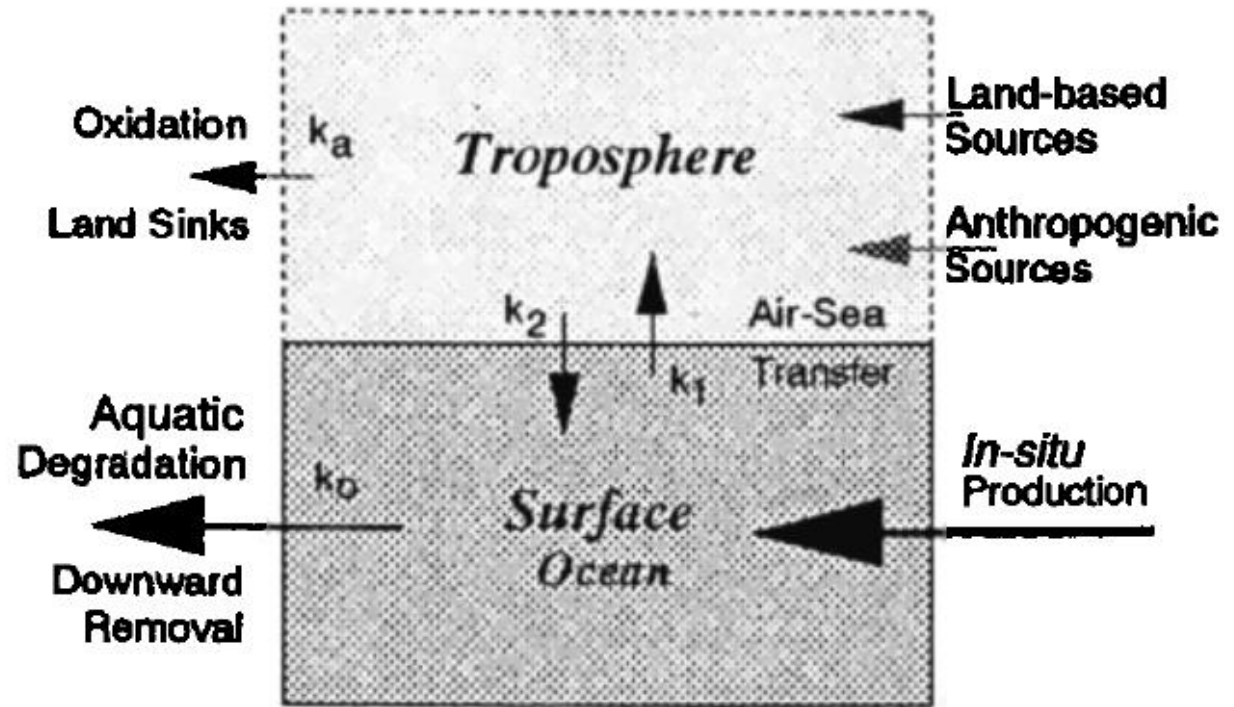
East Pacific Ocean - 1994



Southern Ocean - 1996



# Simple Model



$$\frac{dM_a}{dt} = S_A + F_l + k_1 M_o - k_2 M_a - k_a M_a$$

$$\frac{dM_o}{dt} = P - k_1 M_o + k_2 M_a - k_o M_o$$

# Simple lifetime equation

$$\frac{dM_a}{dt} = S_A + F_l + \frac{K_w A}{H} (p_w - p_a) - k_{OH} p_a M - k_s p_a M$$

$$\frac{dM_o}{dt} = P_o + \frac{K_w A}{H} (p_w - p_a) - k_d M_o - \frac{K_{ed}}{z} M_o$$

$$k_1 = \frac{K_w}{z}$$

$$k_2 = \frac{K_w A}{HM}$$

$$k_o = k_d + \frac{K_{ed}}{z}$$

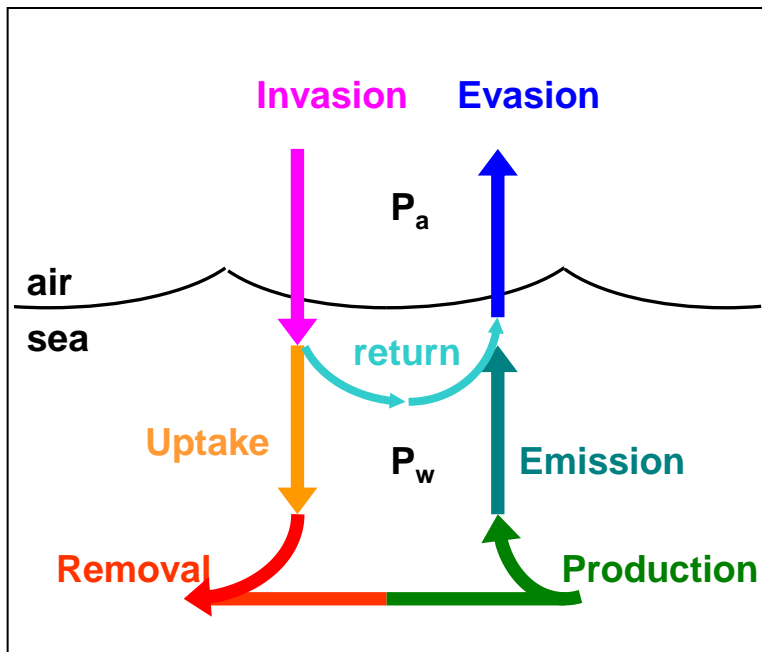
In situ degradation

$$R = \frac{k_o}{(k_1 + k_o)}$$

Return to atmosphere

$$\frac{1}{\tau_o} = k_2 R$$

Partial atmospheric lifetime  
with respect to the ocean



## Terms for Calculations

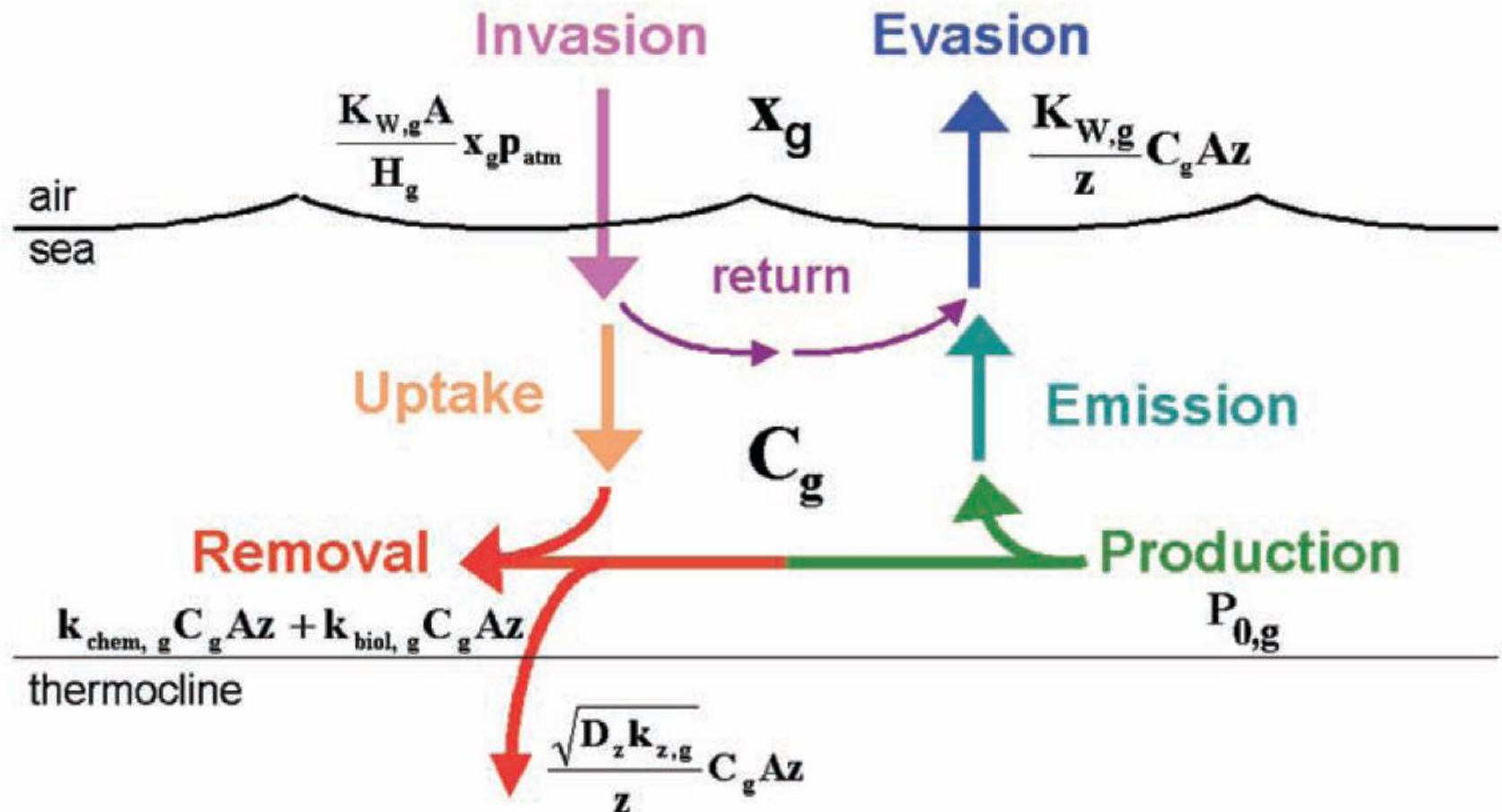
| Term        | Formula                               | Equivalent   |
|-------------|---------------------------------------|--|
| Invasion    | $\frac{K_w A p_{ga}}{H_g}$            | Uptake + Return  |
| Evasion     | $\frac{K_w A p_{gw}}{H_g}$            | Emission + Return  |
| Return Flux | $(1-R) \frac{K_w A p_{ga}}{H_g}$      | Invasion - Uptake  |
| Uptake      | $\frac{n_{tr} p_{ga}}{.95 \tau_o}$    | Evasion - Emission<br>Invasion - Return                    |
| Production  | $P$                                   |  |
| Removal     | $\frac{k_d A z p_{gw}}{H_g}$          | Prod - Emis + Upt  |
| Emission    | $(1 - R) P$                           | Evasion - Return<br>Net Flux - Uptake                      |
| Net Flux    | $\frac{K_w A (p_{gw} - p_{ga})}{H_g}$ | Prod. - Removal<br>Evasion - Invasion<br>Emission - Uptake |

## • Measure

- Partial Pressure difference
- SST
- Salinity
- Windspeed

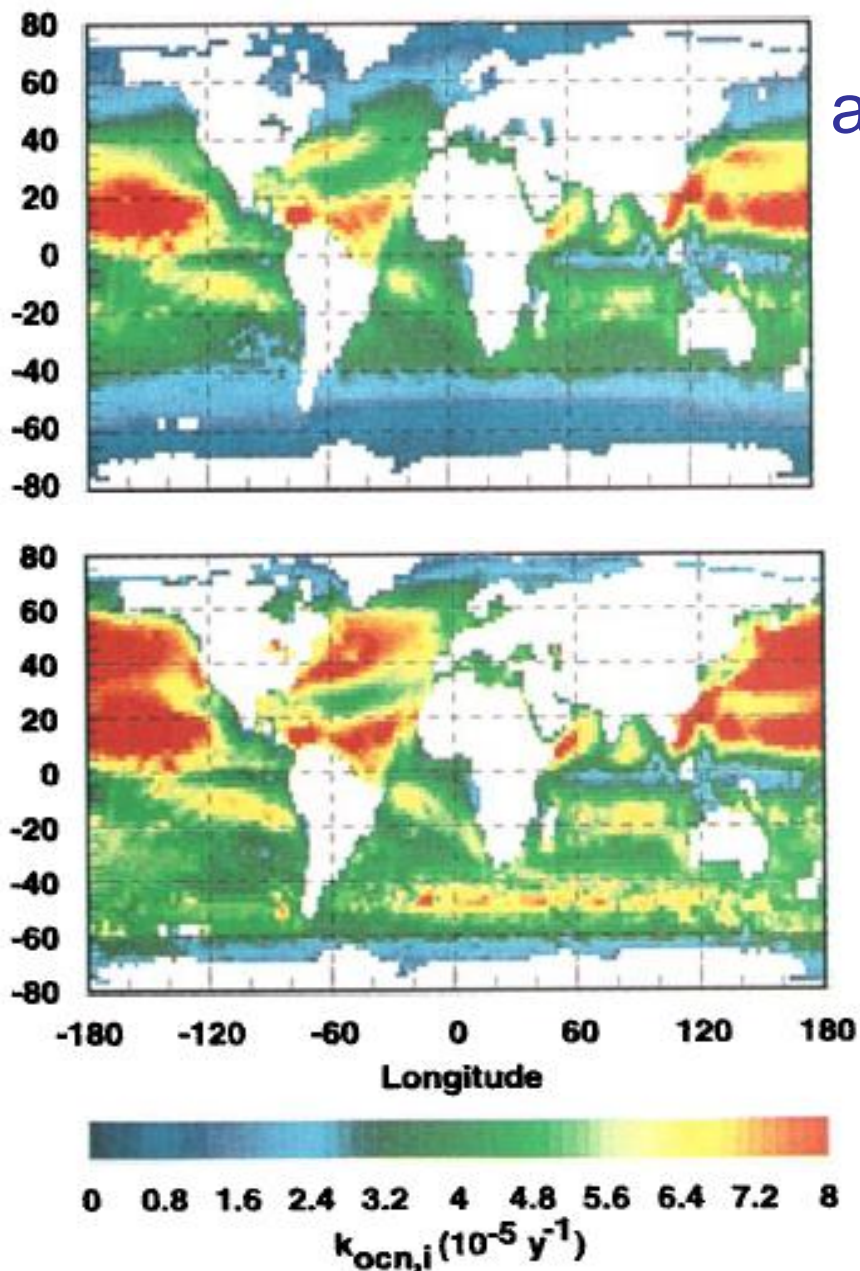
## • Calculate

- Solubility
- Diffusivity
- Viscosity
- Air-sea exchange coefficient



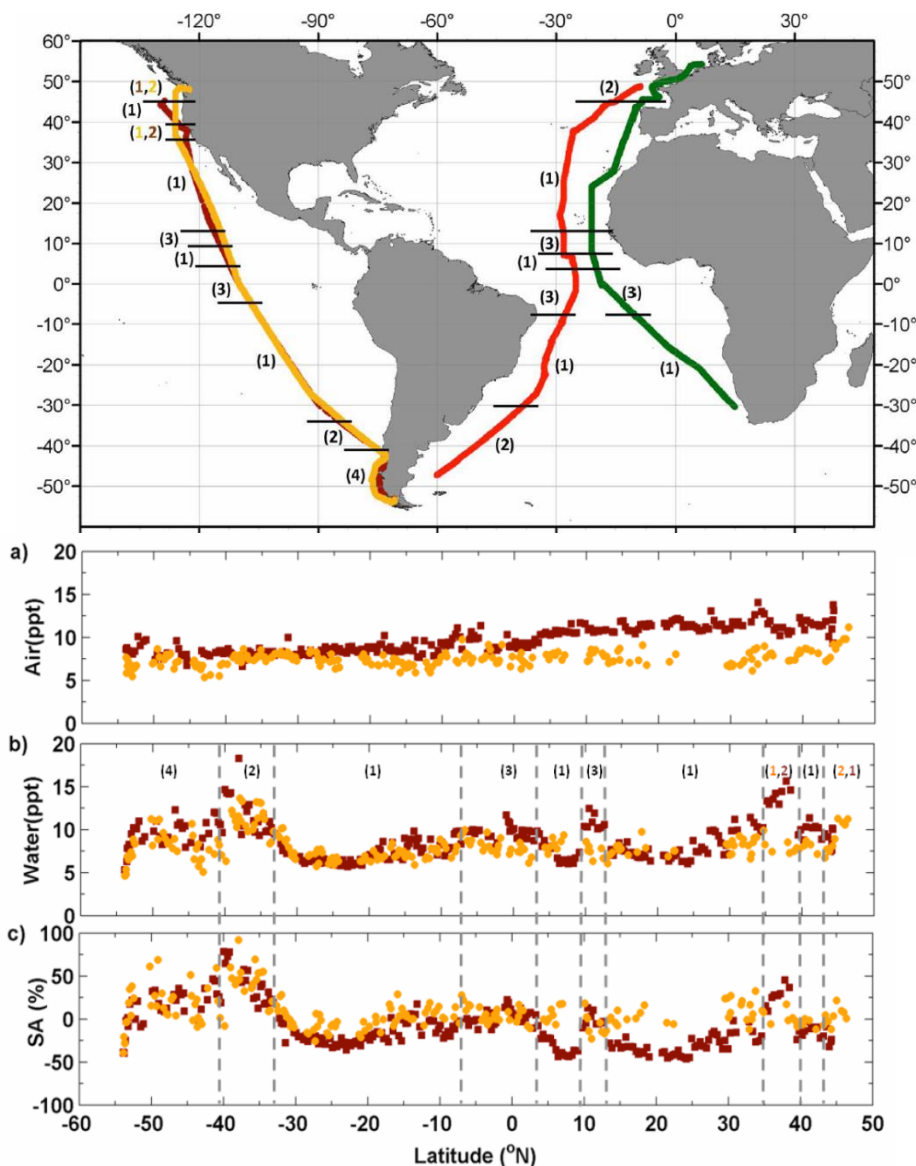


# Influence of changes on atmospheric lifetime of $\text{CH}_3\text{Br}$

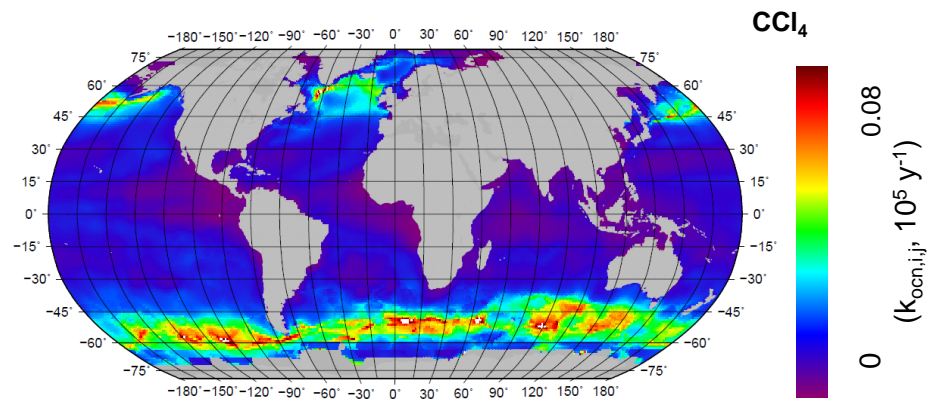
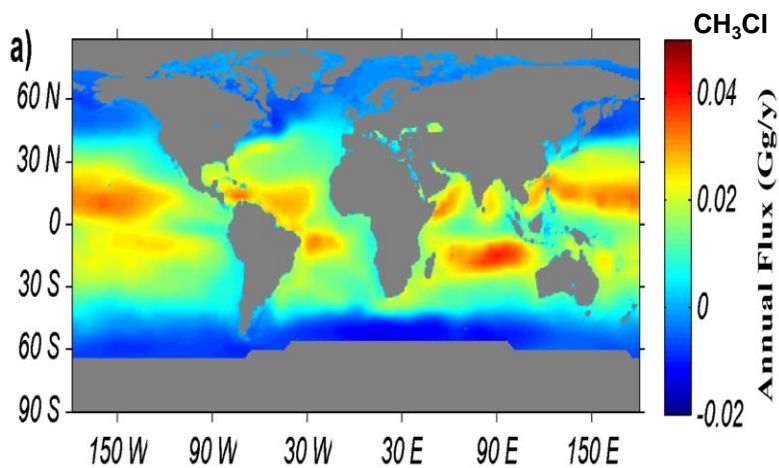
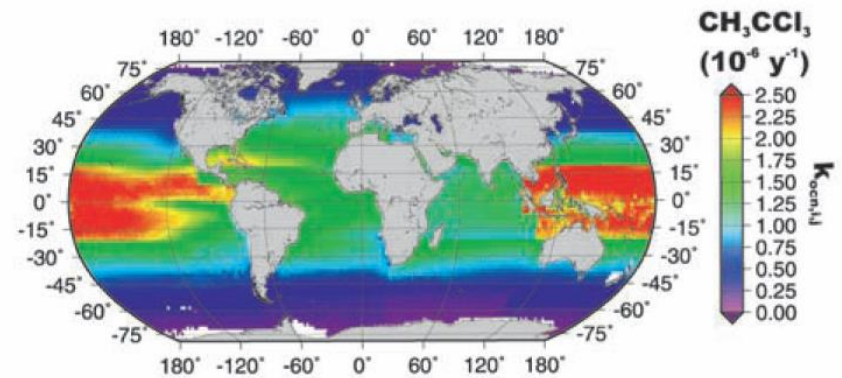
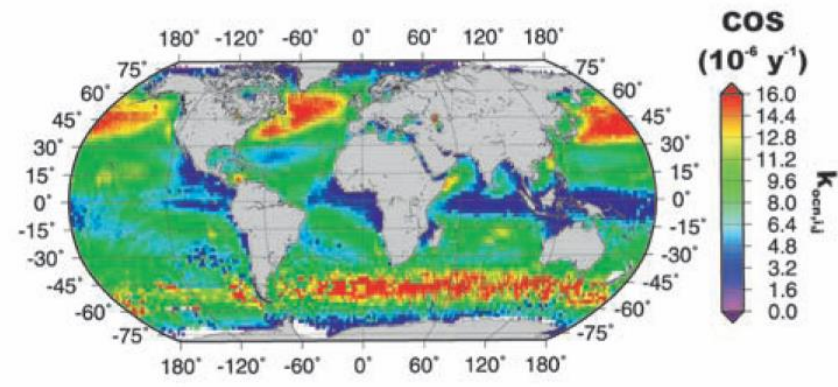
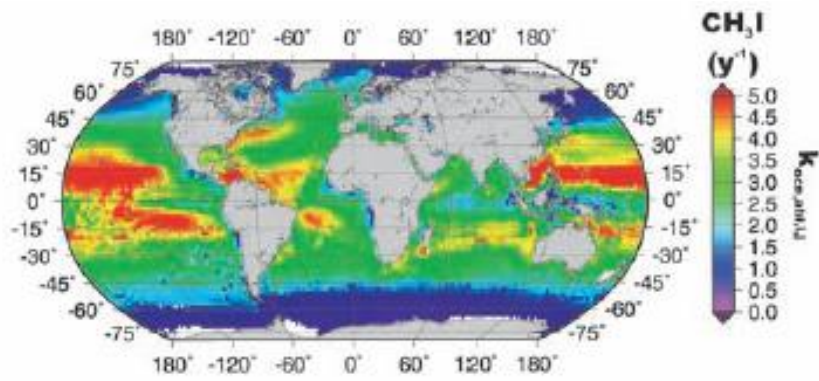


- No Ocean (and no soils)
  - $\tau(\text{atm})$  2.2 y
- Two-box model
  - $\tau(\text{atm})$  1.2 y
  - Range 0.7-1.8 y
- $1^\circ \times 1^\circ$  model (chem only)
  - $\tau(\text{atm})$  0.8 y
  - Range 0.6-1.4 y
- $1^\circ \times 1^\circ$  model (chem+ bio)
  - $\tau(\text{atm})$  0.7 y
  - Range 0.6-0.9
- 2014 Ozone Assessment
  - $\tau(\text{atm})$  0.80 y
- 2014 with  $W_{(2014)} K_w$ 
  - $\tau(\text{atm})$  0.84 y

# Footnote for CH<sub>3</sub>Br



- Lei Hu et al., 2012, Oceanic Saturation State of CH<sub>3</sub>Br, *Global Biogeochemical Cycles*
- Human Emissions reduced
- Atmospheric amount has dropped from ~11 to 7.5 ppt since 1999.
- Net flux is now positive
  - 0 to 3Gg/y
  - Previously ~ -14Gg/y !! (2002 Scientific Assessment)
- Atm. Lifetime still 0.8 (0.6-0.9) y



...can do this  
for any gas!

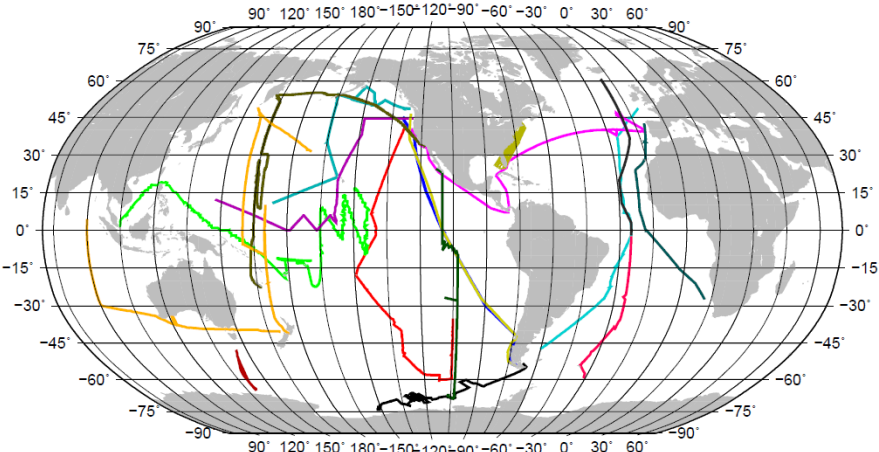


# Our Latest Effort – CCl<sub>4</sub>

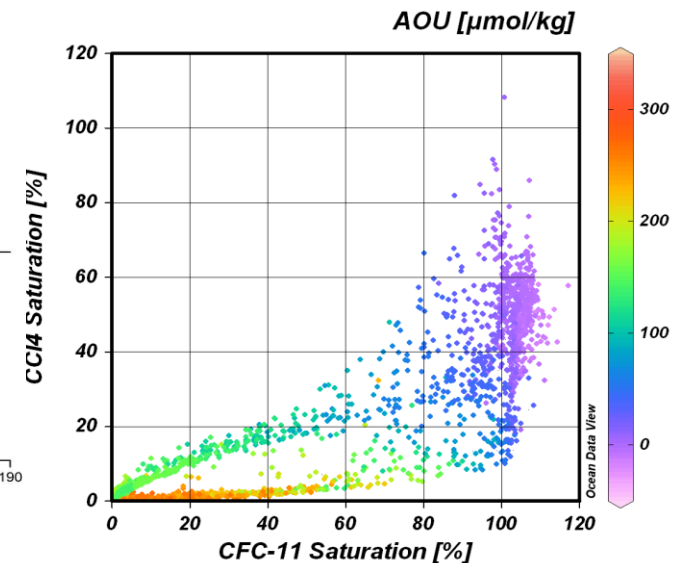
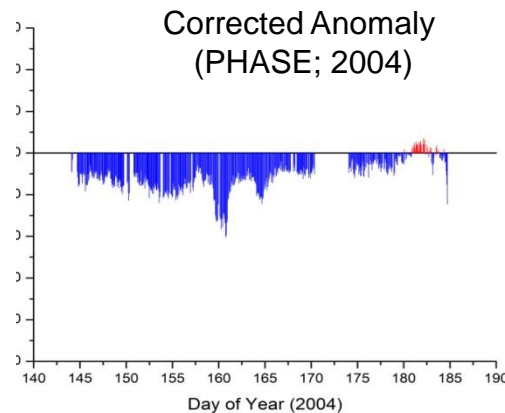
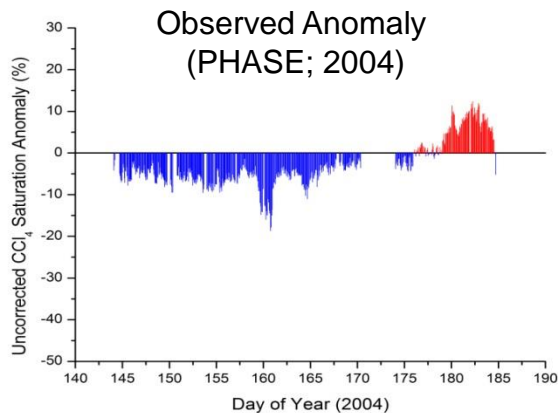
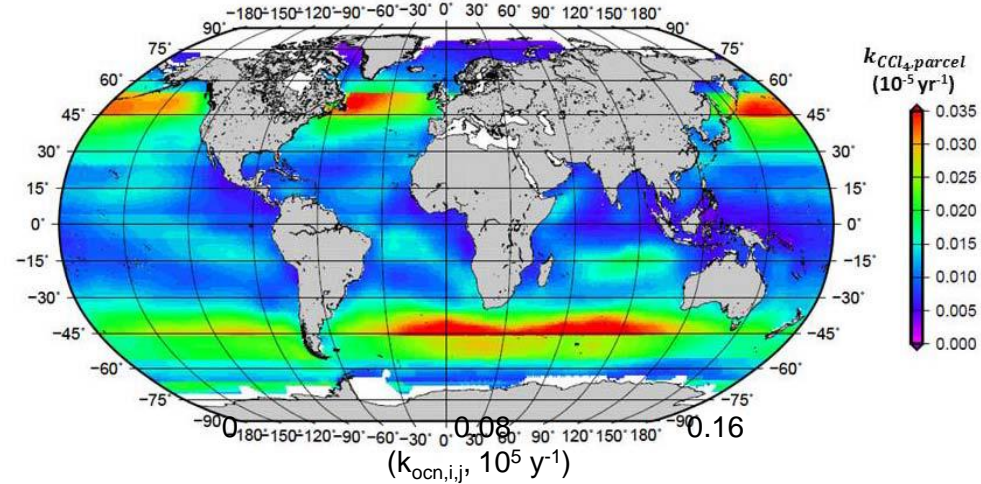


Research cruises

Oceanic Uptake  
Rate Constant



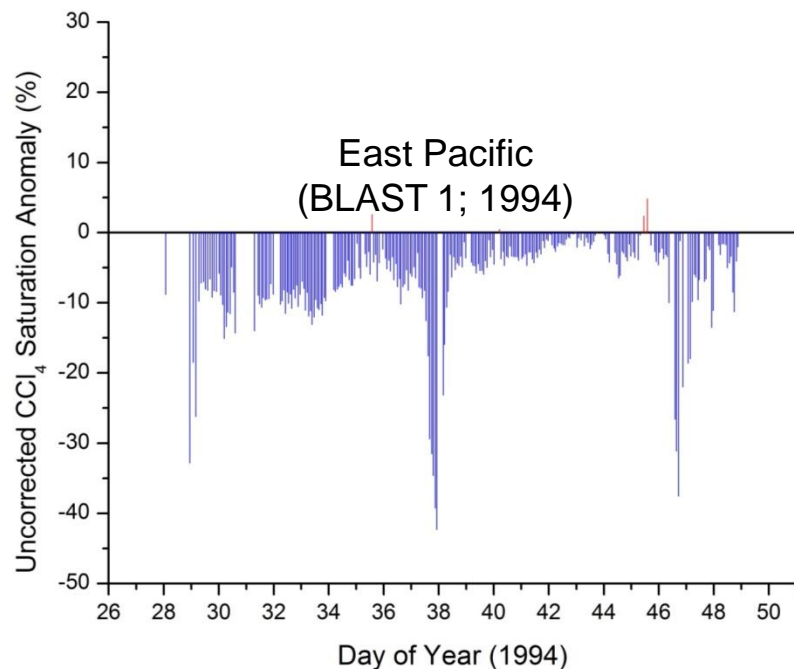
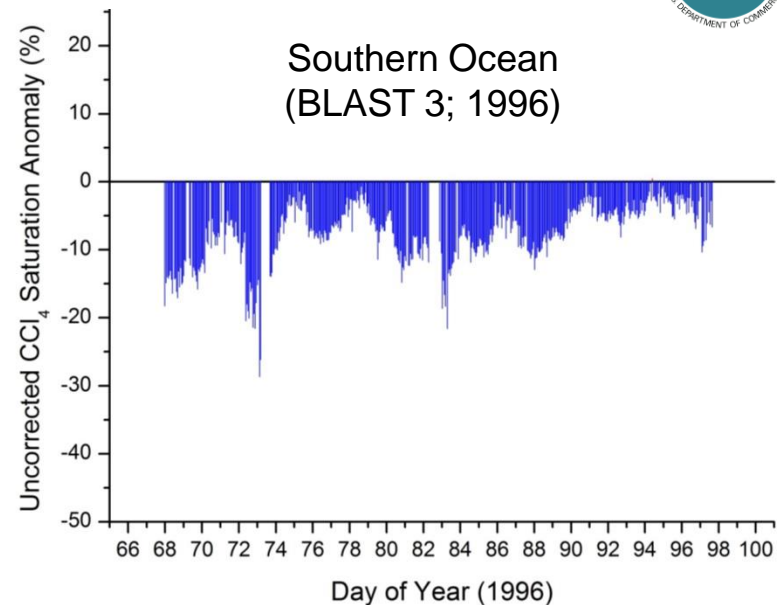
SAGA-2, RITS89, SAGA-3, OAXTC, BLAST1, BLAST2,  
BLAST3, GasEx98, RB9906, CLIVAR01, A16N, A16S,  
PHASE, P18, GOMECC, HalocAST-P, HalocAST-A



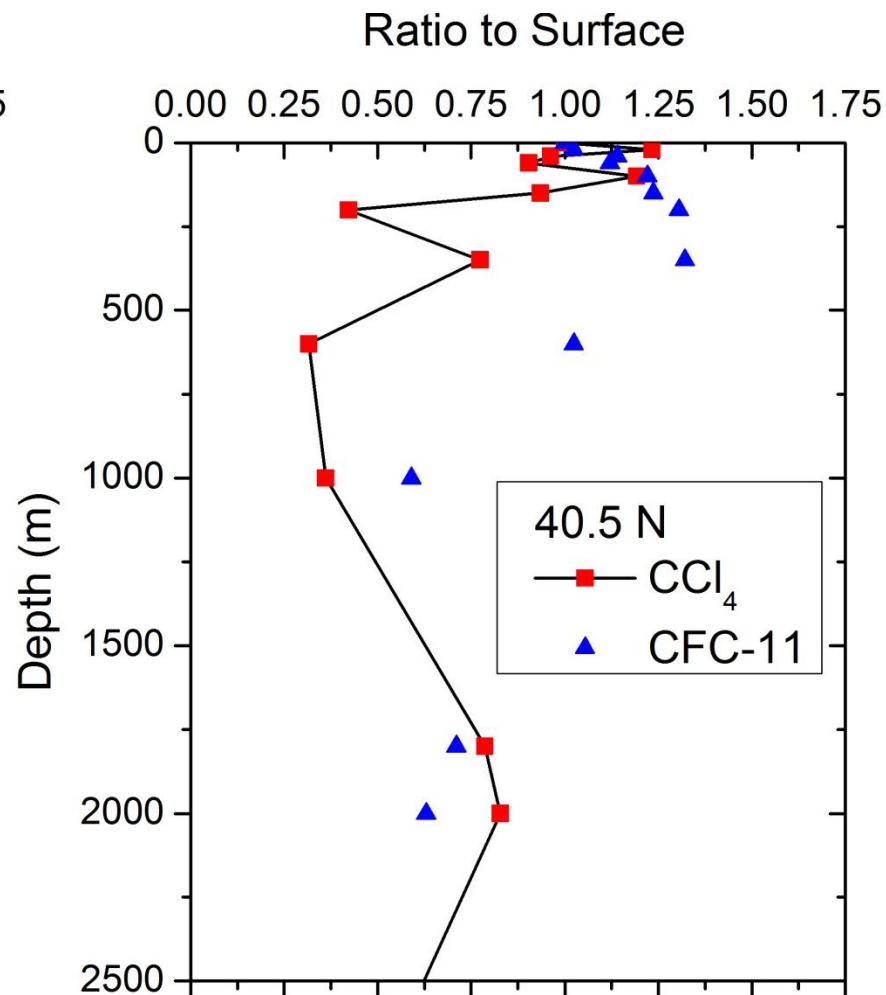
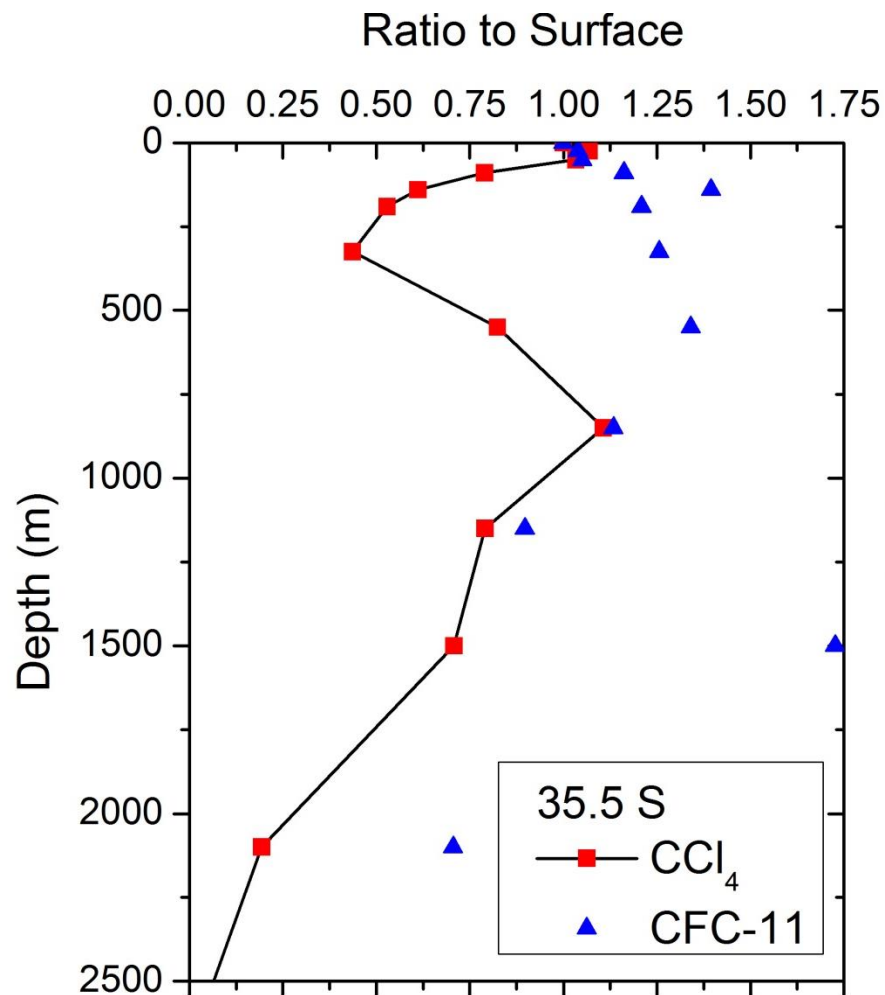


# Biological (?) Consumption of $\text{CCl}_4$ (1987-2011)

- $\text{CCl}_4$  is under-saturated everywhere, all the time
  - No relationship to temperature
  - No significant chemical sink
- But  $\text{CCl}_4$  is consumed at mid-depth ( $\text{O}_2$  minimum)
  - $\text{CCl}_4$  is deficit carried to the surface by mixing and transport
  - But not enough to explain deficit
- Exchange with the atmosphere leads to 18% of atmospheric  $\text{CCl}_4$  being removed irreversibly by the ocean
  - Lowers atmospheric lifetime
  - Lowered ozone-depletion potential
  - Policy-relevant



# CCl<sub>4</sub> at Depth



# How much does $K_w$ change things for $\text{CCl}_4$ ?



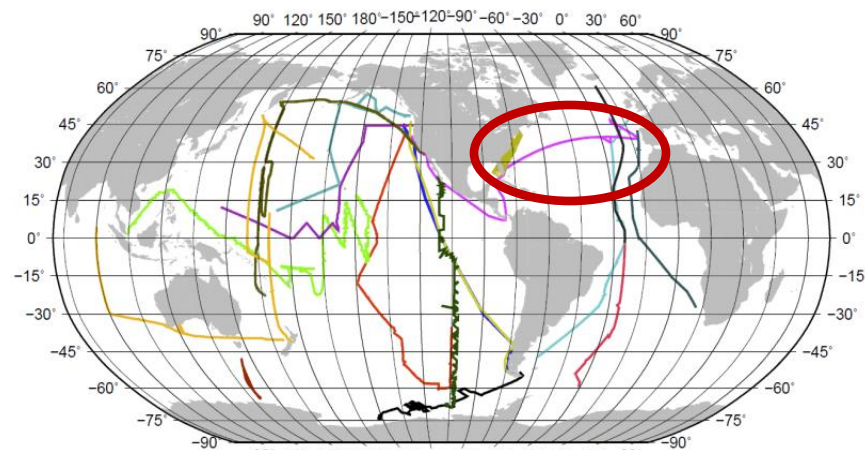
|                                | $W_{(1992)}$ | $W_{(2014)}$ | % change |
|--------------------------------|--------------|--------------|----------|
| $T_{\text{ocn}} \text{ (y)}$   | 139          | 180          | 30%      |
| $T_{\text{strat}} \text{ (y)}$ | 44           | 44           | —        |
| $T_{\text{soil}} \text{ (y)}$  | 375          | 375          | —        |
| $T_{\text{total}} \text{ (y)}$ | 30.6         | 32.3         | 6%       |

# Do these gases have utility for evaluating $K_w$ ?

| CFC-11         |       | Kw        |     | sd   | cv%    |
|----------------|-------|-----------|-----|------|--------|
|                |       | 4.9 m/d   |     | 0.8  | 16%    |
|                |       | 20.4 cm/h |     | 3.2  |        |
| Property       | Mean  | sd        | n   | se   | cv(se) |
| hz (m)         | 50    |           |     |      | 5.0%   |
| dt (d)         | 46.26 | 1.92      | 116 | 0.18 | 0.4%   |
| Kappa (%/deg)  | 4.48  |           |     | 0.07 | 1.6%   |
| dT (deg)       | 3.93  | 0.96      | 114 | 0.09 | 2.3%   |
| Del(g) end (%) | 3.81  | 2.36      | 39  | 0.38 | 9.9%   |
| Del(g) avg (%) | 3.04  | 1.98      | 37  | 0.32 | 10.7%  |

| CFC-12         |       | Kw        |     | sd   | cv%    |
|----------------|-------|-----------|-----|------|--------|
|                |       | 6.3 m/d   |     | 1.4  | 23%    |
|                |       | 26.2 cm/h |     | 6.0  |        |
| Property       | Mean  | sd        | n   | se   | cv(se) |
| hz (m)         | 50    |           |     |      | 5.0%   |
| dt (d)         | 46.26 | 1.92      | 116 | 0.18 | 0.4%   |
| Kappa (%/deg)  | 3.95  |           |     | 0.06 | 1.6%   |
| dT (deg)       | 3.93  | 0.96      | 114 | 0.09 | 2.3%   |
| Del(g) end (%) | 2.83  | 2.35      | 33  | 0.41 | 14.5%  |
| Del(g) avg (%) | 2.18  | 2.11      | 33  | 0.37 | 16.9%  |

| CH3CCI3             |       | Kw        |     | sd   | cv%    |
|---------------------|-------|-----------|-----|------|--------|
| Thermal Diseq alone |       | 13.7 m/d  |     | 7.0  | 51.3%  |
|                     |       | 57.1 cm/h |     | 29.3 |        |
| with Hydrolysis     |       | 4.4 m/d   |     |      |        |
|                     |       | 18.2 cm/h |     |      |        |
| Property            | Mean  | sd        | n   | se   | cv(se) |
| hz (m)              | 50    |           |     |      | 5.0%   |
| dt (d)              | 46.26 | 1.92      | 116 | 0.18 | 0.4%   |
| Kappa (%/deg)       | 4.04  |           |     | 0.06 | 1.6%   |
| dT (deg)            | 3.93  | 0.96      | 114 | 0.09 | 2.3%   |
| Del(g) end (%)      | 3.03  | 3.12      | 38  | 0.51 | 16.7%  |
| Del(g) avg (%)      | 1.01  | 2.93      | 36  | 0.49 | 48.2%  |



Variability =  $\pm 19\%$

[Wanninkhof (2014) uncert =  $\pm 20\%$ ]



