

# The OceanFlux Greenhouse Gases project

Aims to improve the quantification  
of air-sea exchanges of greenhouse  
gases, of prime importance  
in the climate system.

## Parameterizing the bubble- mediated air-sea flux of a non- ideal gas, DMS

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## Bubble Effects

- Enhancement in gas transfer rate
- Equilibrium super-saturation condition [Woolf 1997]
- Physical-Chemical effects
  - .....What are bubble effects for non-ideal gases such as DMS
    - chemical properties may be useful in predicting bubble effects for non-ideal gases such as methylated sulfur species
    - Couple with air-sea bubble model

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## Addressing solubility in a bubbly surface ocean

In senescent conditions

$$C_{i,w} \approx \gamma_{i,w}^{-1} V_i^{-1}$$

At high wind speeds (>12 m/s)

$$C_{i,mix} \approx \gamma_{i,mix}^{-1} V_i^{-1} \approx (1-x_b)C_{i,w} + (x_b) K_H C_{i,w} + K_{b-film} C_{i,w}$$

$$\approx [(1-x_b)\gamma_{i,w}^{-1} + (x_b)\gamma_{i,b}^{-1} + \Phi_B \gamma_{i,b-film}^{-1}] V_i^{-1}$$

or

Where  $\Phi_B = f(x_b) =$  bubble surface area per  $m^2$

$\gamma_{i,j} =$  activity coefficient of substrate  $i$  in phase  $j$

$V_i =$  molar volume of substrate  $i$  in  $cm^3 mol^{-1}$

$x_b =$  volume fraction of bubbles in sample volume



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## Henry's Law Constant, $K_{iH}$

$K_{iH}$  = vapor pressure/solubility

$$= p_i / C_{i,w}$$

$$= \gamma_{iw} \rho_{iL} V_w$$

$K_i = K_{iH} / RT$  (dimensionless)

$$K_{i,mix} = K_i / (1 + (C_{mix} / C_w) \Phi_B)$$



$K_{b-film}$

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In terms of flux

$$\text{Flux} = k_i \Delta C_{AW}$$

$$k_i = [1/(k_a K_{i,mix}) + 1/k_w]^{-1}$$

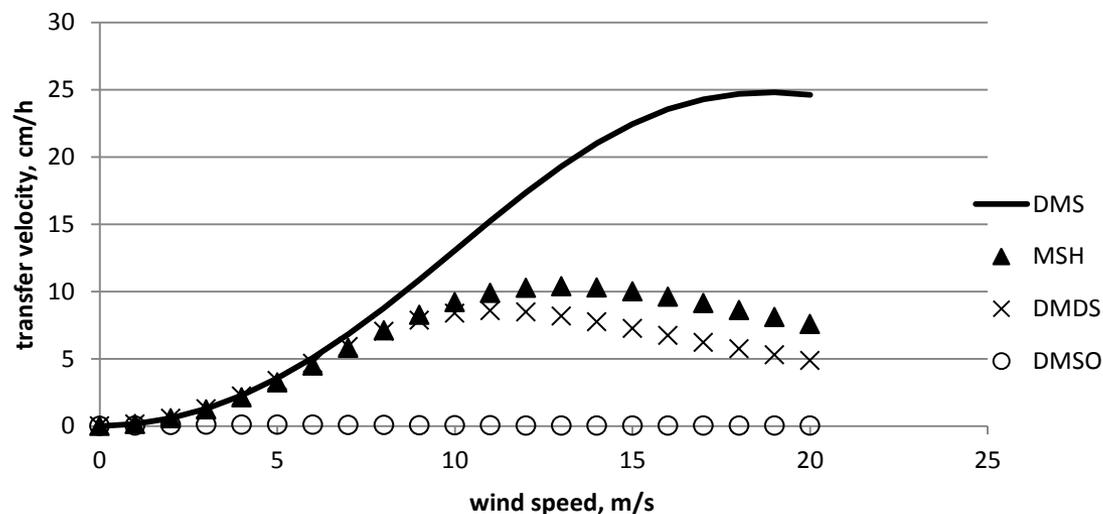
$$\text{where } K_{i,mix} = K_i / (1 + (C_{mix}/C_w) \Phi_B)$$

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## Predicted influence of wind speed on transfer velocities of surface active gases

| Compound | Chemical Formula                             | MWt , g/mol | Solubility, (mol/L) | K, (dimensionless, air/water) |
|----------|--|-------------|---------------------|-------------------------------|
| MSH      | CH <sub>3</sub> SH                           | 48.11       | 0.195               | 0.0115                        |
| DMS      | C <sub>2</sub> H <sub>6</sub> S              | 62.13       | 0.355               | 0.08                          |
| DMSO     | C <sub>2</sub> H <sub>6</sub> SO             | 74.12       | miscible            | 5.25 x 10 <sup>-7</sup>       |
| DMDS     | C <sub>2</sub> H <sub>6</sub> S <sub>2</sub> | 94.20       | 0.0363              | 0.0428                        |

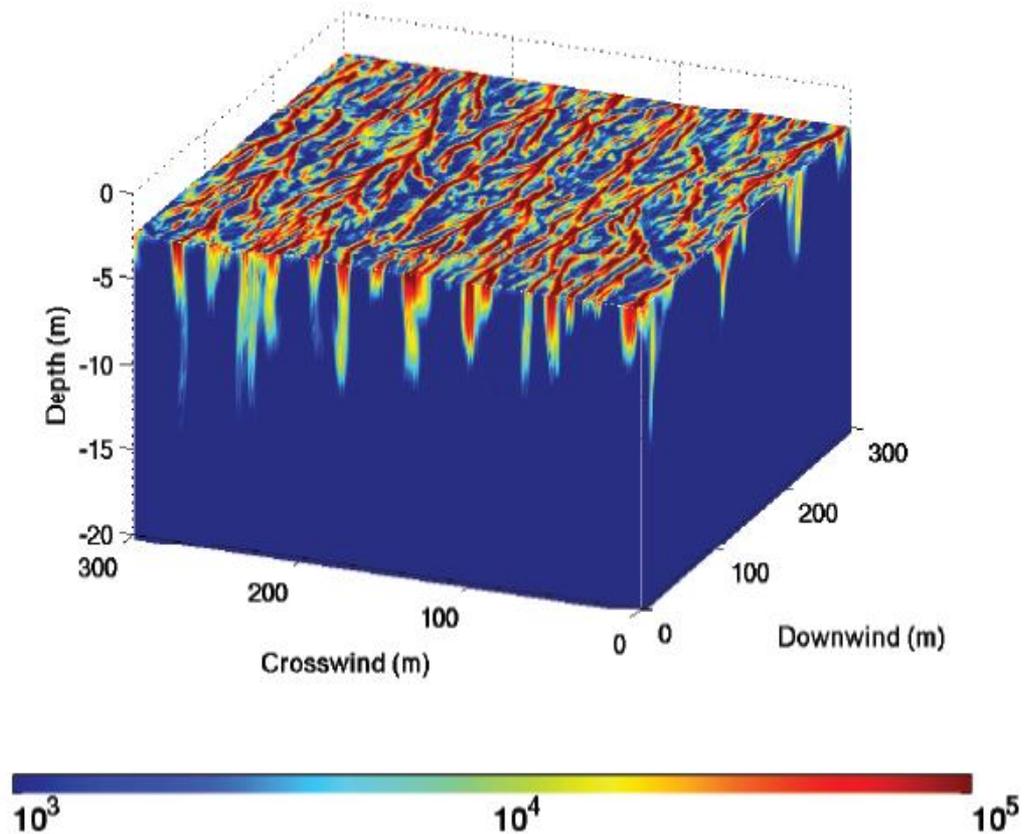


From Vlahos et al (2011), *Gas Transfer at Water Surfaces*, 2010

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Liang et al. 2011, 2012, 2013

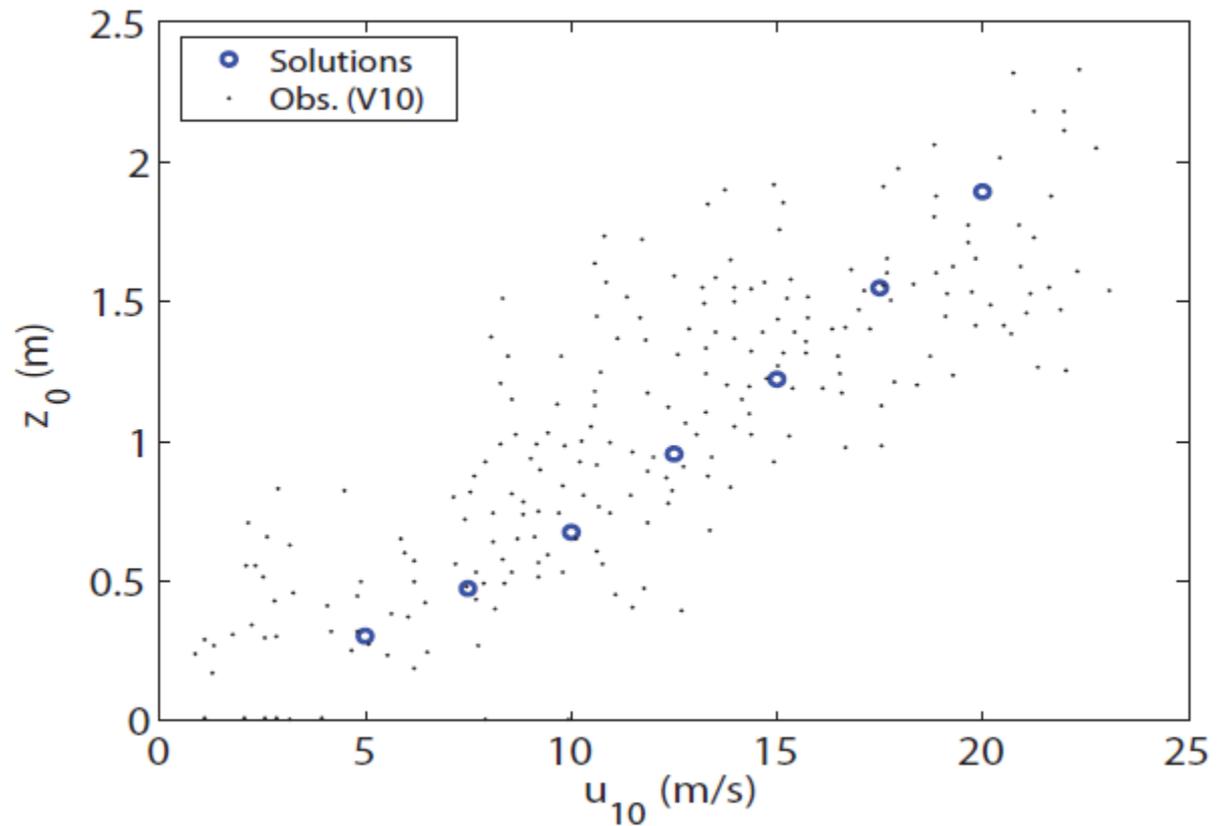


simulated total bubble number density  
(number per m<sup>3</sup>) at  $u_{10} = 17.5$  m/s.

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Liang et al. 2013



Relationship between bubble e-folding depth  $z_0$  and wind speed  $u_{10}$ . (Black dots: observations by *Vagle et al.* [2010]; Blue circles: bubbly flow solutions assuming waves in equilibrium with local winds.)

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## This model

- bubbles of sizes **>35 microns** are included
- **fraction** of different gases change with time
- considers the effect of ambient **pressure change** on bubble size change
- readily **coupled** to upper ocean dynamic model
- Bubble injection is **constrained** by laboratory and *in situ* observations

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## Total Gas Flux

$$F = F_{sfc} + F_p + F_c + SP \frac{C_{mix}}{C_w} \int \frac{d\Phi}{dt} dz$$

- $F_{sfc}$  is the gas flux through the ocean surface
- $F_p$  is the gas flux through large bubbles that partially dissolve
- $F_c$  the gas flux through small bubbles that completely dissolve
- the last term is an effective gas flux due to the change of bubble surface area assuming a proportionate fraction of molecules attach

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## Estimated from Liang Model

$$F = k_{sfc} (S_{eff} P - C) + k_b [S_{eff} P(1 + \Delta_{overpressure}) - C] + F_c + SP \frac{C_{mix}}{C_w} \int \frac{d\Phi}{dt} dz$$

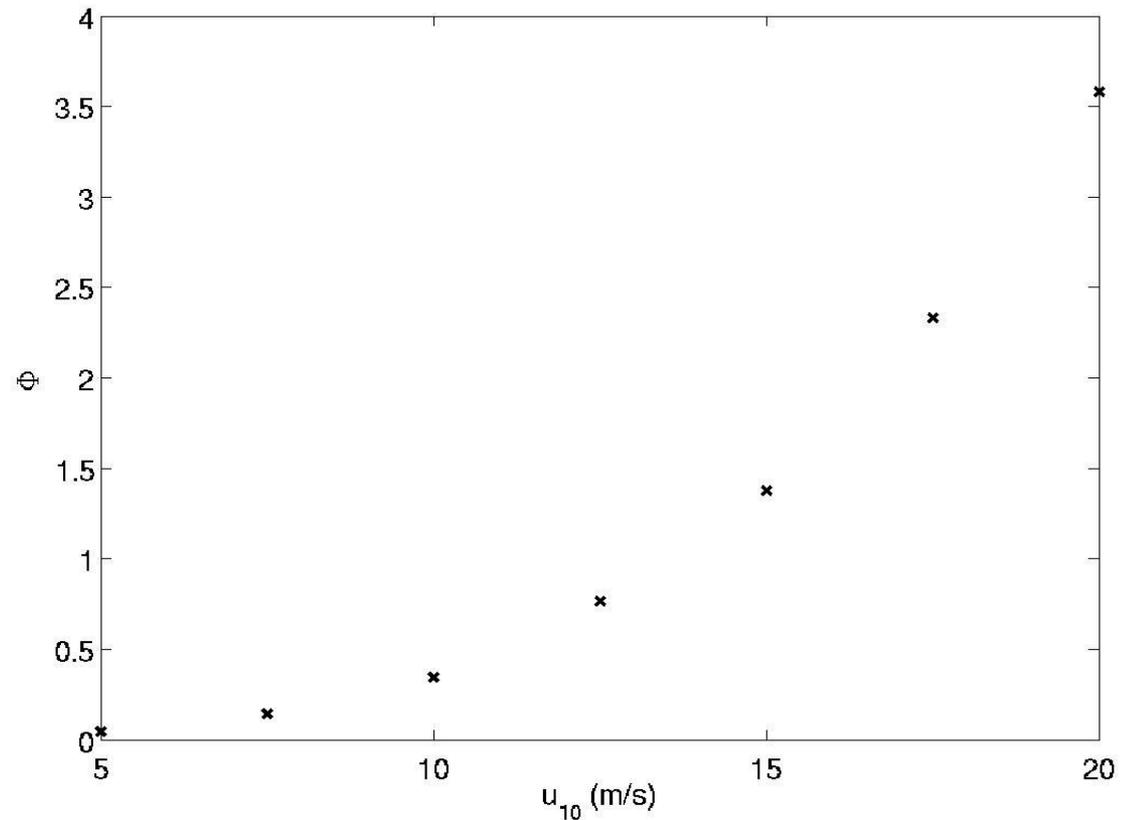
$$F = k_{tot} [S_{eff} P(1 + \Delta) - C]$$

$$\Delta = \frac{k_b}{k_T} \Delta_{overpressure} + \frac{F_c + SP \frac{C_{mix}}{C_w} \int \frac{d\Phi}{dt} dz}{k_T SP \left( 1 + \frac{C_{mix}}{C_w} \Phi \right)}$$

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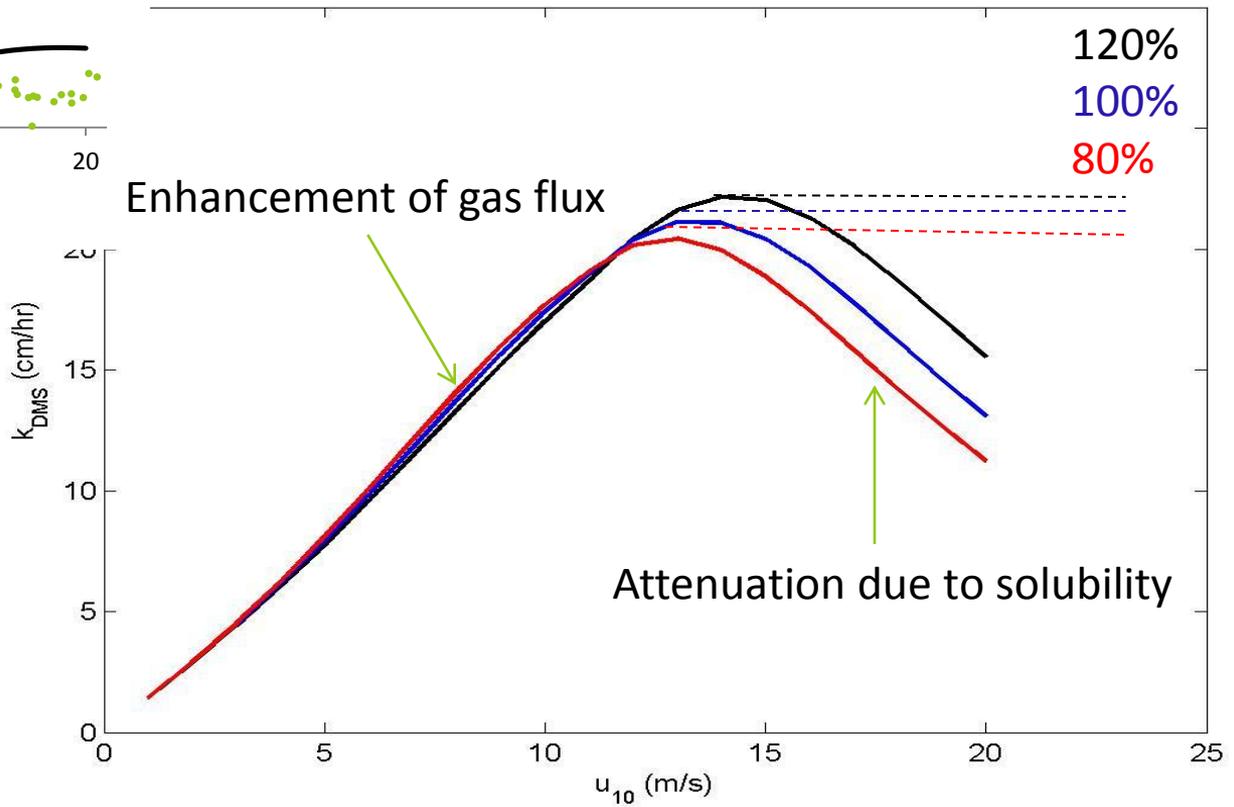
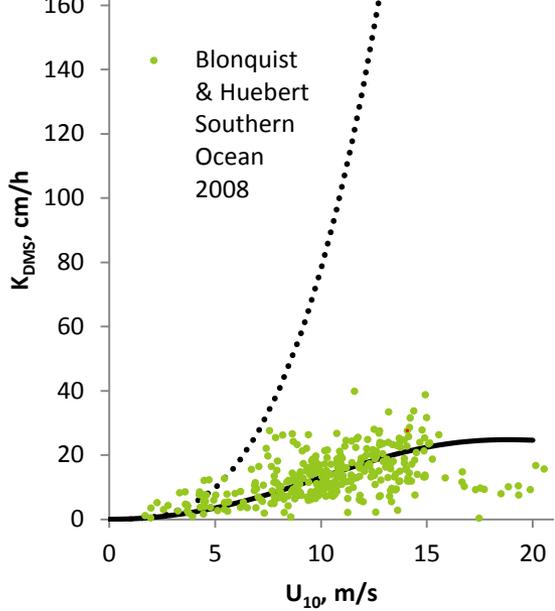
Aims to improve the quantification of air-sea exchanges of greenhouse gases, of prime importance in the climate system.

## Relationship between $\Phi$ and wind speed



- (1) the total injected bubble amount in the model is based on observations, which are limited, and a source of possible uncertainty;  $\Phi_B = 0.40 (U/10)^3$
- (2) the formula for  $k_w$  in VM09 is at the lower end of observed values  $\Phi_B = 0.090 (U/10)^3$

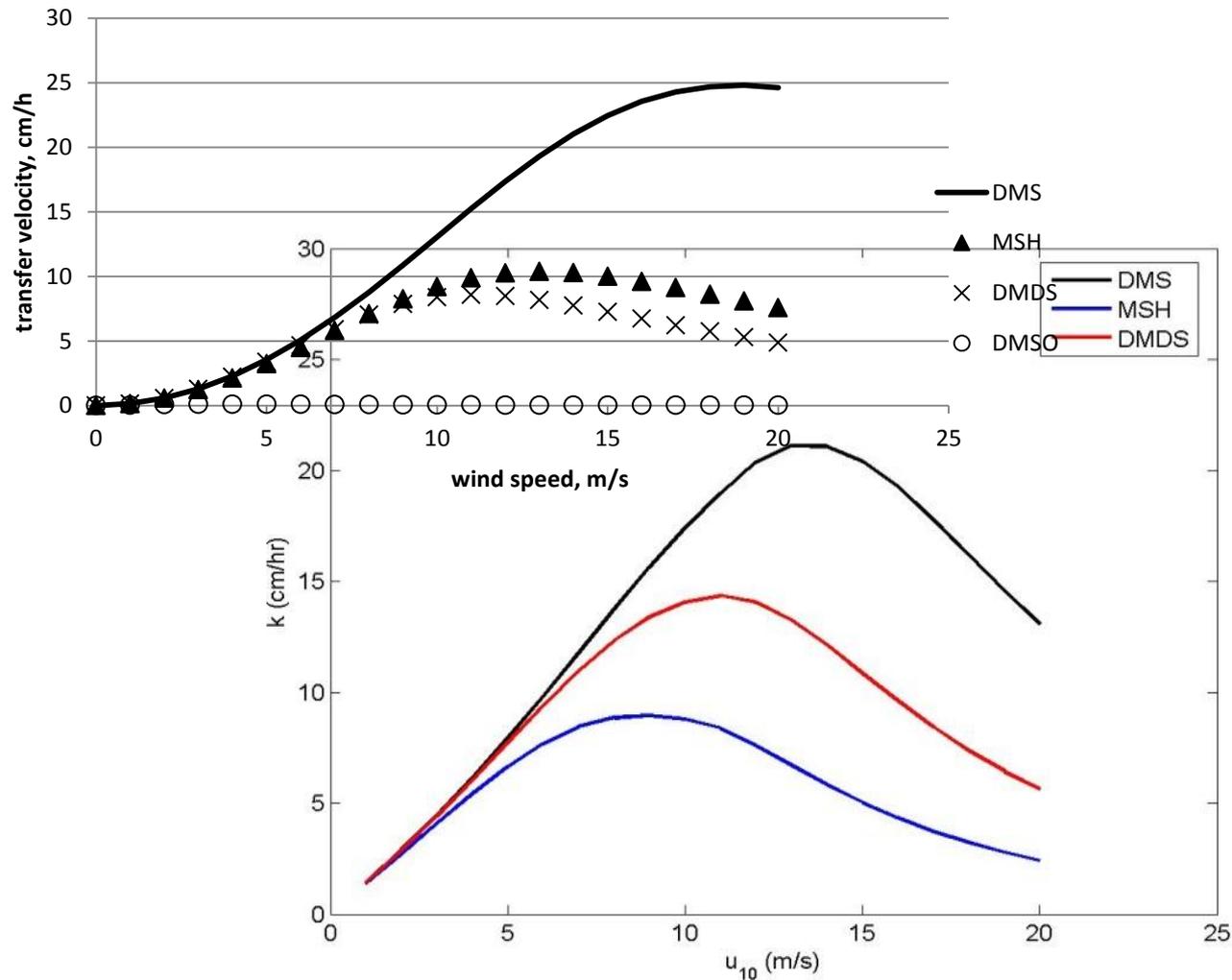
# Sensitivity test on injected bubble amount



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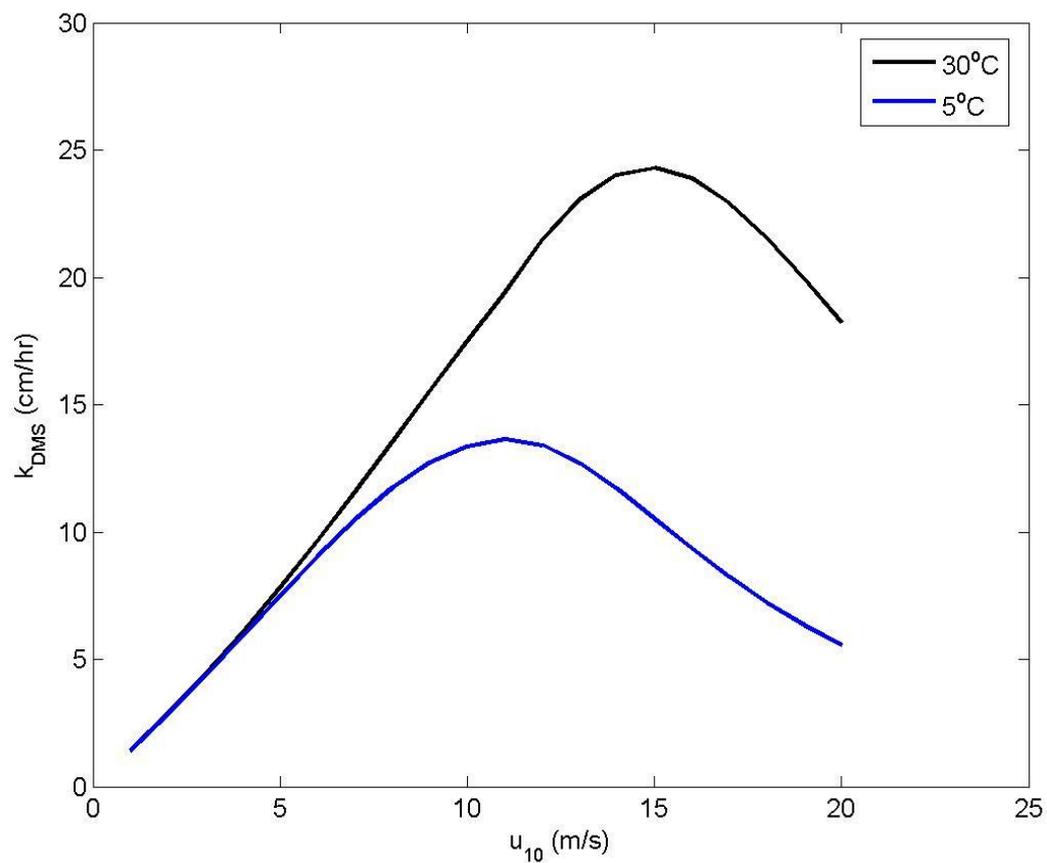
## Gas flux profile for organic gases



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## Temperature effects



- DMS and derivatives expected to be moderately surface active
- At high winds ( $> 12$  m/s ) the surface ocean is significantly altered
  - the effective activity coefficient and therefore the fugacity of surface active compounds is expected to decrease
  - The result is higher relative water-side concentrations in the bubble influenced layer and lower transfer velocities

## Current Results

- Air-sea flux of DMS may be overestimated using traditional empirical models
- DMS air-sea flux reaches upper thresholds governed by a piston **velocity near 25 cm/h** and the corresponding concentration gradient.
- This sets upper DMS air-sea transfer rates at **120  $\mu\text{Mm}^{-2}\text{d}^{-1}$**  based on observed field concentrations. (DMS  $C_w$ 's range from <1 to 20 nM (Lee et al 2010; Zang et al, 2008))

## Current Results

- Sensitivity analysis on nitrogen and oxygen
- Coupling to Lagrangian model
  - Equilibration of fast dissolving gases in partially dissolved bubbles will be explicitly resolved.
  - Bubble size distribution will be better resolved.
  - Add equilibration effect to the model  
Woolf [1993], Keeling [1993]
- Modeling of other of biogenic gases

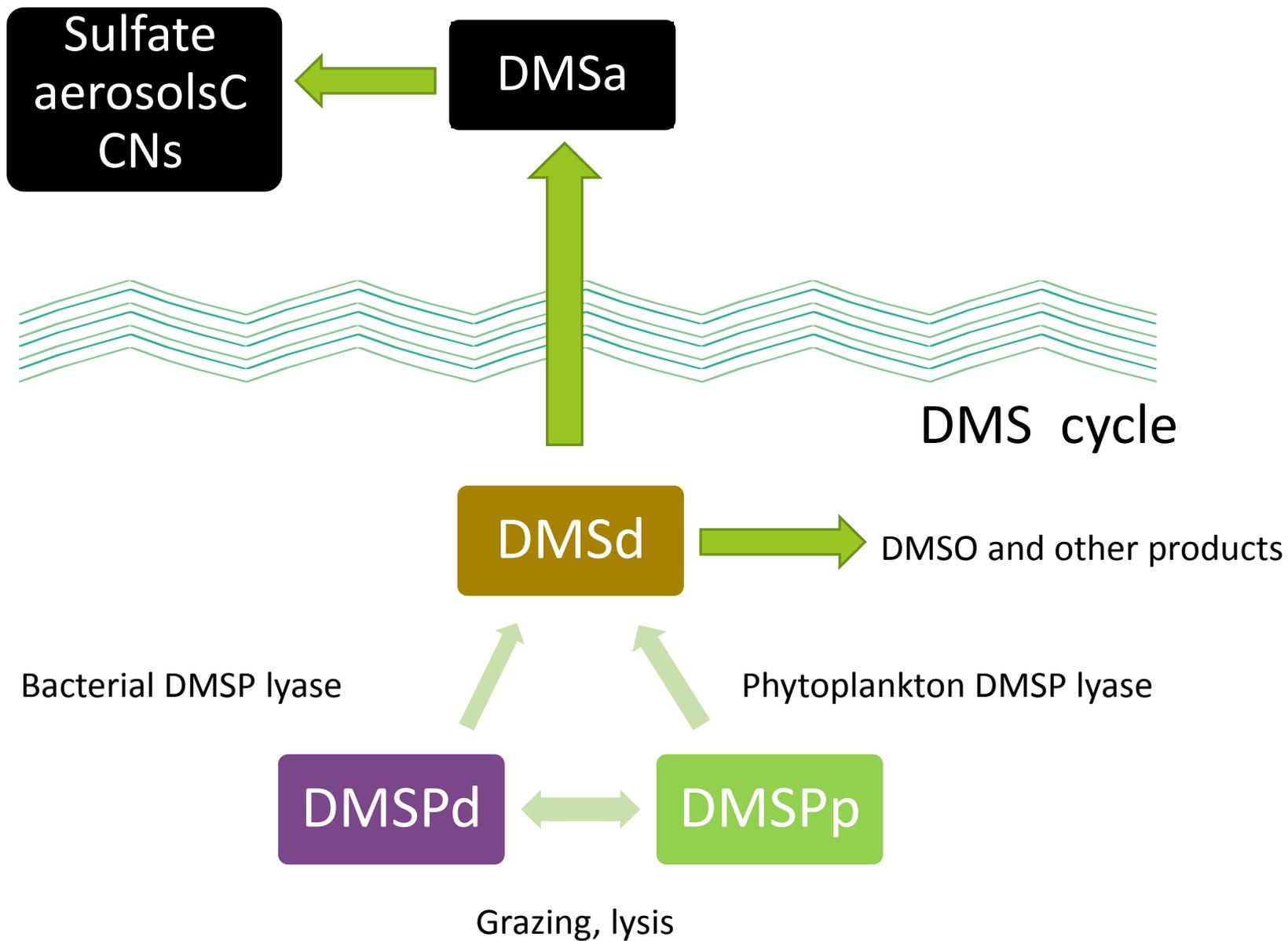
Next Steps...



That's all for now...

- DMS  $C_w$ 's range from <1 to 20 nM (Lee et al 2010; Zang et al, 2008)
- DMSP p + d, 9.22 (2.85-19.73) and 17.50 (4.33-36.09)nM (Zang et al, 2008)
- Photodegradation rates 0.006 to 0.028 nM/(Wm<sup>-2</sup>) (Miles et al., 2009)
- DMS degradation rates 0.04 to 0.66 d<sup>-1</sup>(Kieber et al., 1996)
- DMS production rates 0.7 to 0.9 nMd<sup>-1</sup> (Bailey et al 2008)

What we know

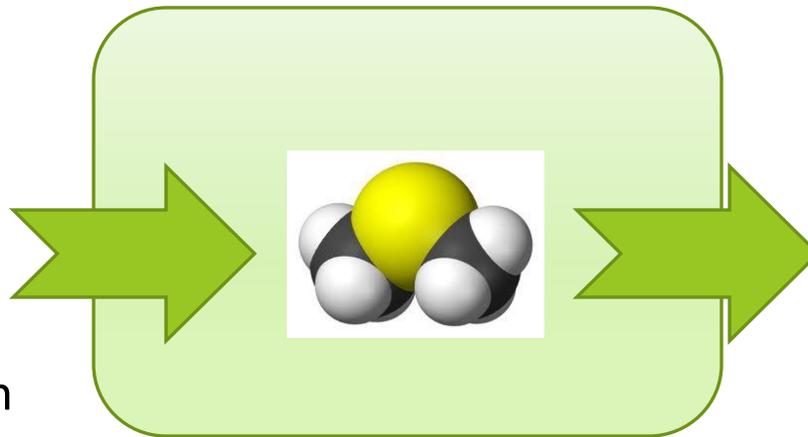




# DMS box model

## Sources:

- ✓ phytoplankton
- ✓ DMSP<sub>d</sub> breakdown
- ✓ advection

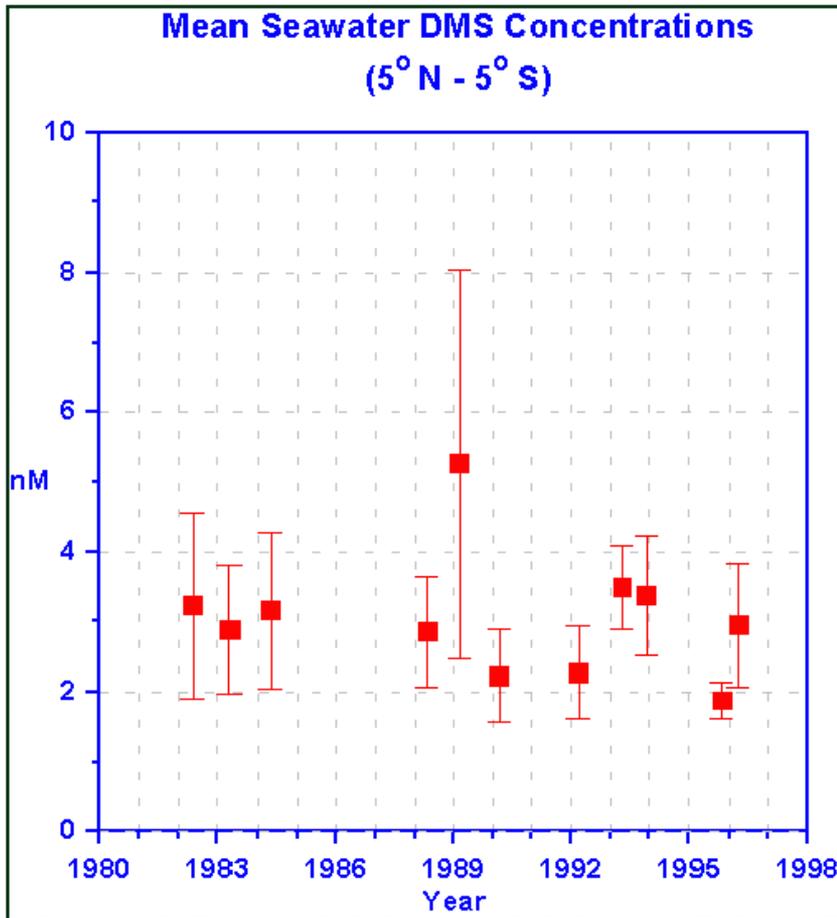


## Sinks:

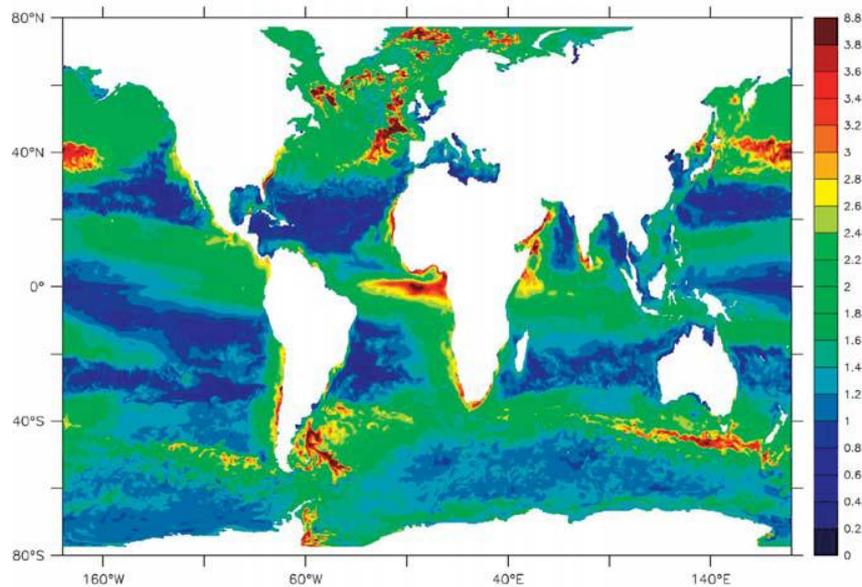
- ✓ Air-sea exchange
- ✓ Biodegradation
- ✓ Photodegradation
- ✓ advection

Equatorial Pacific (PMEL)- 15°N to 15°S DMS is relatively constant both seasonally and interannually ( $2.7 \pm 0.7$  nM (Bates and Quinn, *Geophys. Res. Lett.*, 24:861-864, 1997)).

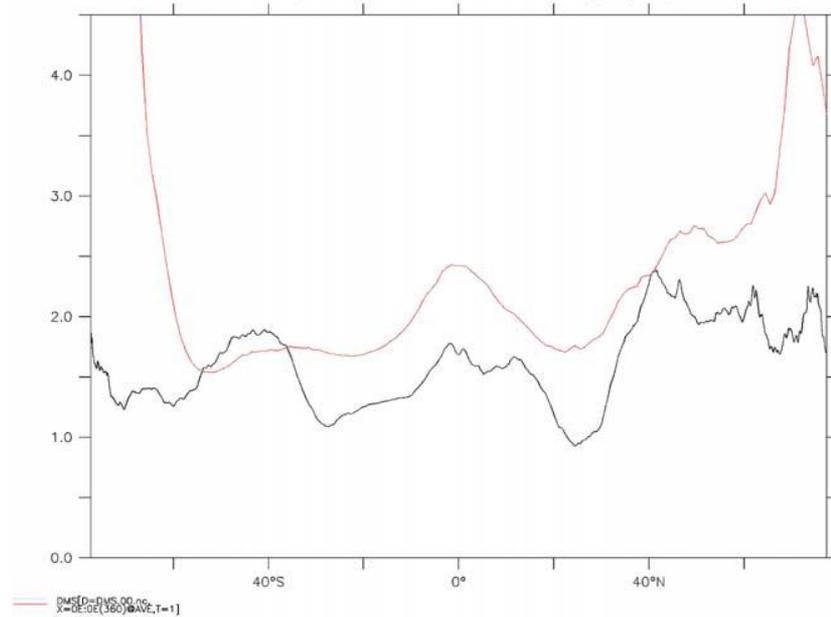
## Observations



1a – Annual surface DMS (nmol)

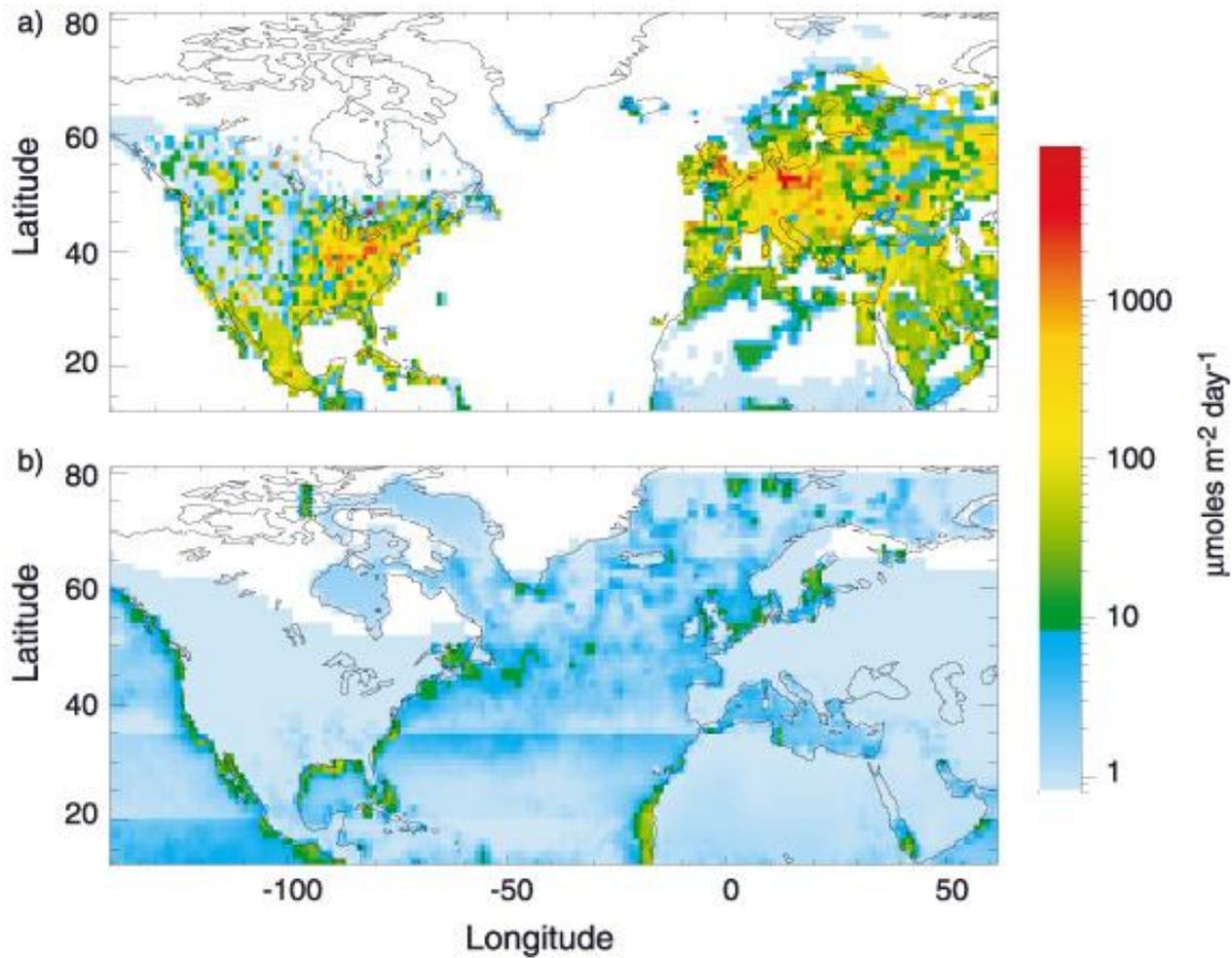


1b – Annual zonal average surface DMS (nmol)  
POP (black) and Kettle Climatology (red)



## Surface ocean DMS

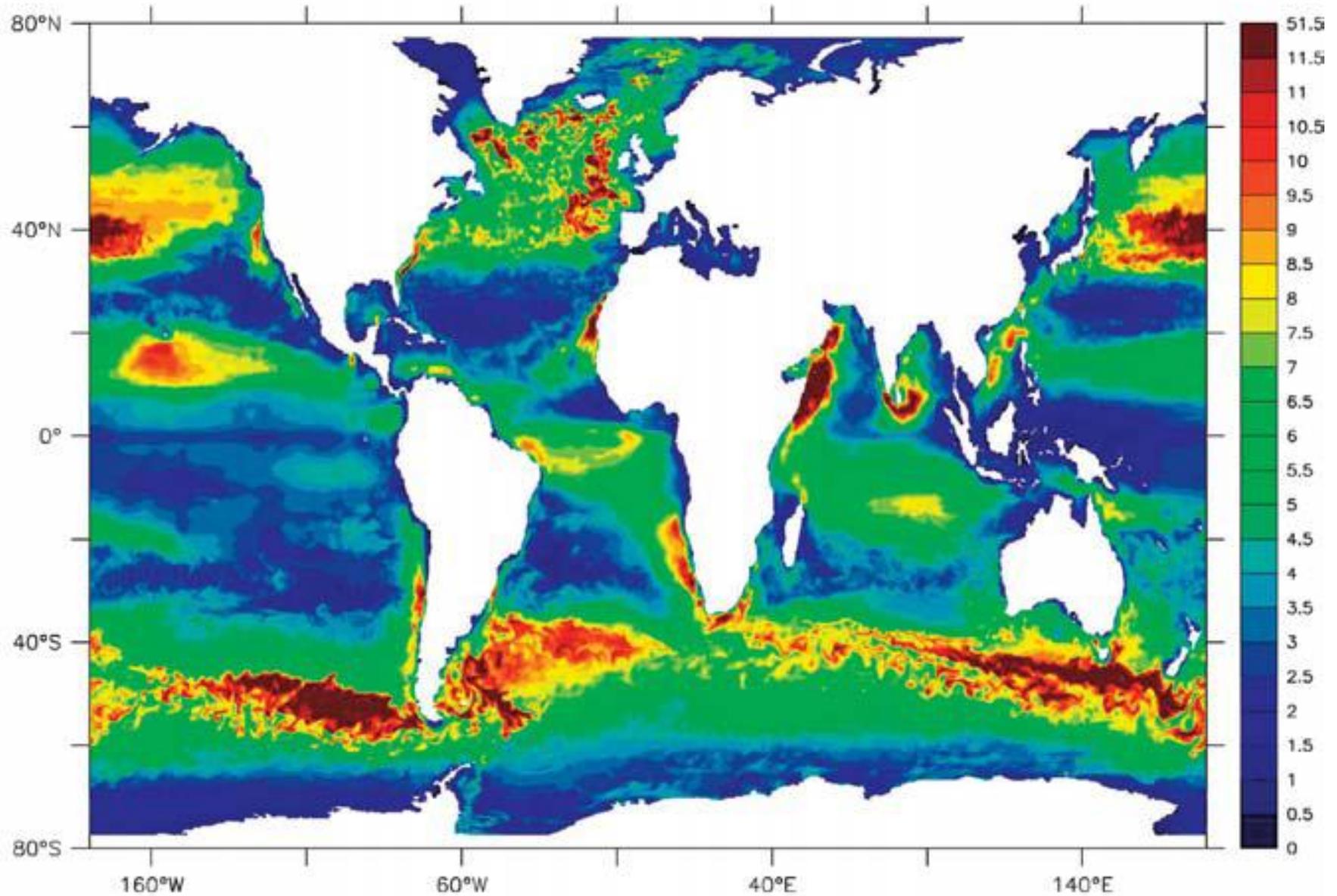
Earth Interactions Volume 8 (2004)



DMS fluxes

Bates et al. [1992]

7a – Annual DMS air–sea flux ( $\text{mmol}/\text{m}^2/\text{day}$ )



7b – Annual zonal average DMS air–sea flux ( $\text{mmol}/\text{m}^2/\text{day}$ )