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Remote sensing algorithms for sea surface CO₂ and CO₂ flux

a Swedish National Spaceboard project
focusing the Baltic Sea

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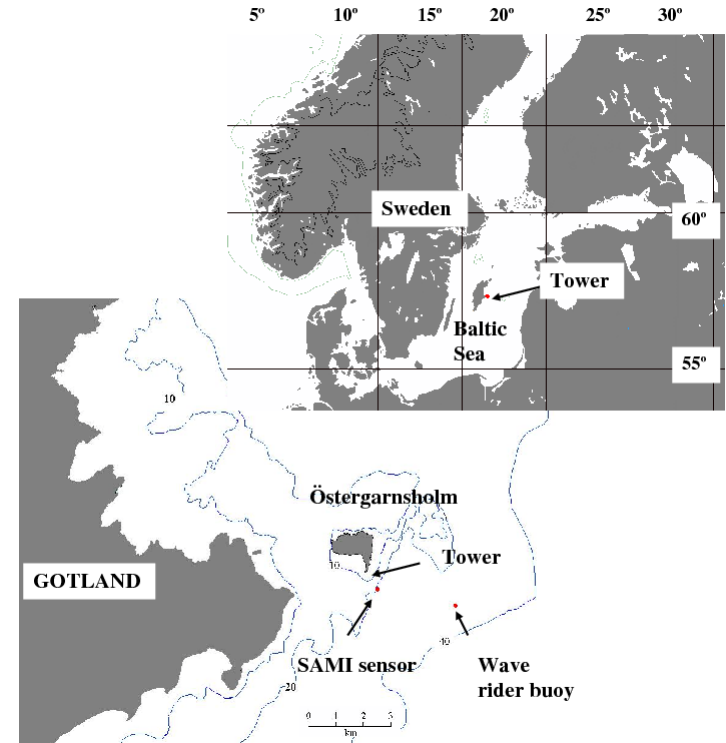
Baltic Earth

Earth System Science for the Baltic Sea Region



From a global perspective: Why Baltic Sea?

1. Representative of coastal regions
(contribution of marginal seas and shelf areas to the ocean carbon cycle is disproportionately large).
2. Highly variable
 - river runoff
 - upwelling
 - atmosphere more variable
(stratification, sea-breeze circulation, coastal jets etc).





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Do we need to consider water-side convection when calculating air-sea gas transfer?

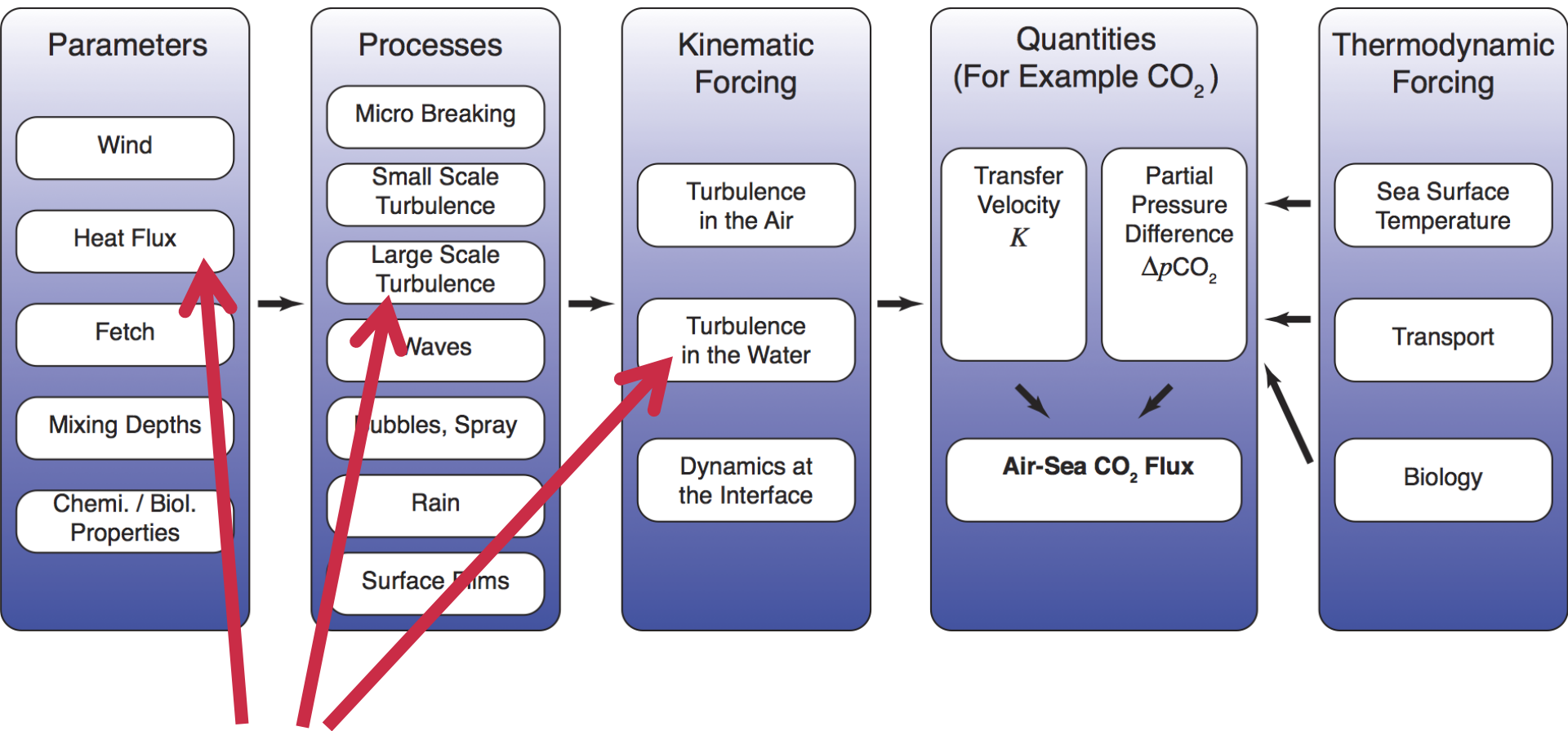
Transfer velocity for different gases.

Anna Rutgersson

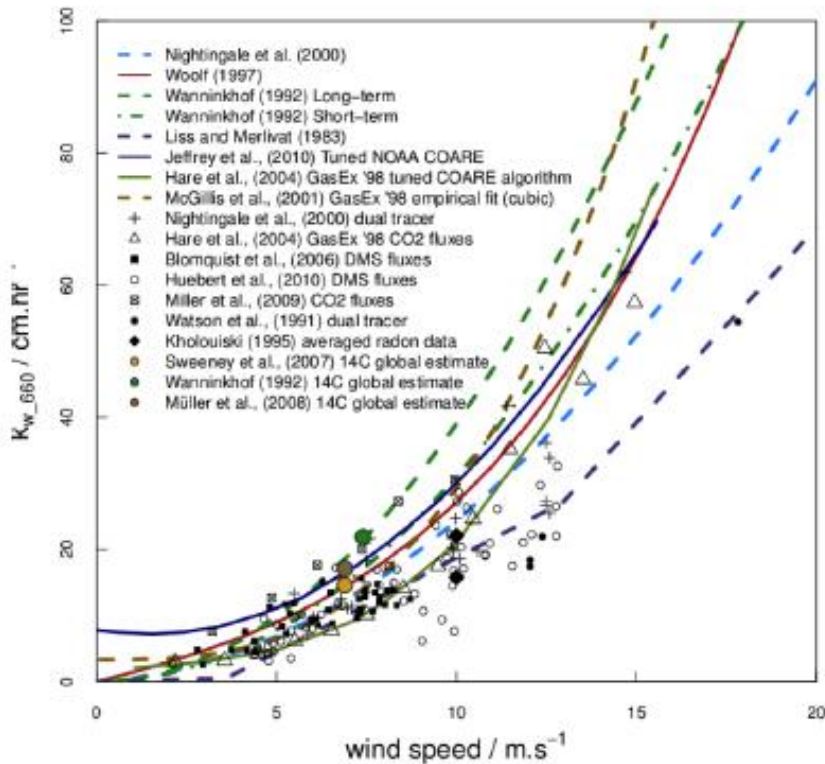
also Erik Sahlée, Eva Podgrajsek, Maria Norman, Andreas Andersson



What controls the exchange?



Wind-speed dependent transfer velocity



$$F_{CO_2} = \overline{w'c'} \quad (\text{turbulent vertical flux})$$

$$F_{CO_2} = kK_0(pCO_{2w} - pCO_{2atm})$$

$$k \propto U^x$$

$$x = 1, 2, 3, \dots$$

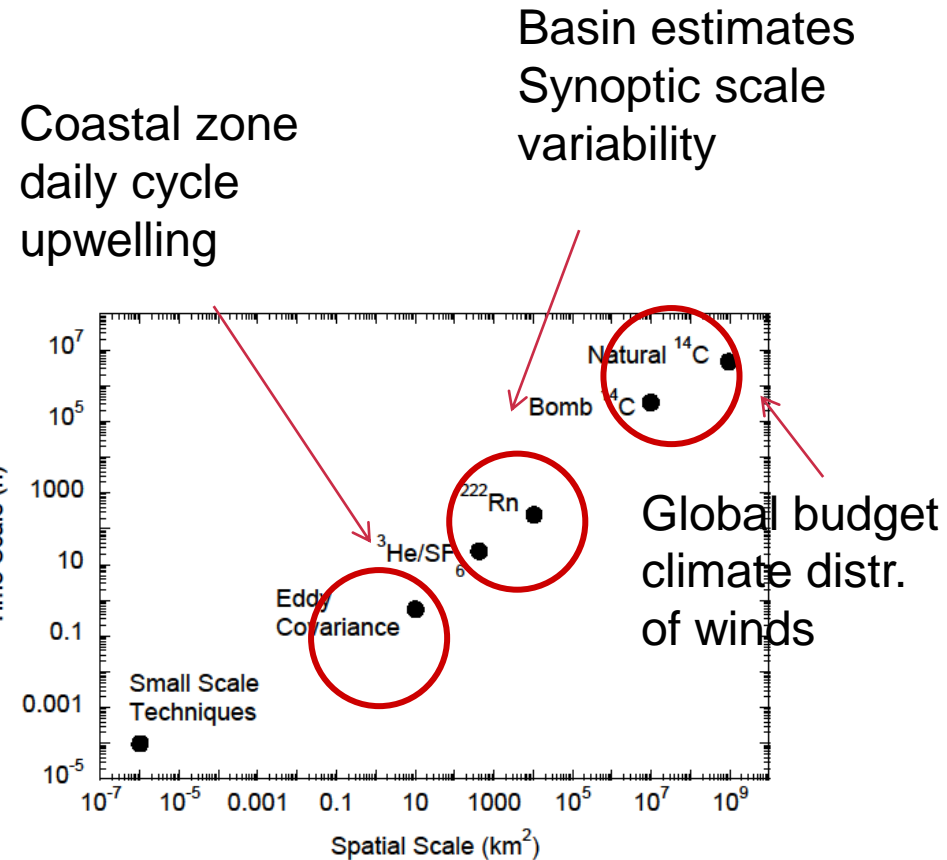
Fig. 4 A comparison of different tracer windspeed relationships of the transfer velocity k .

Time scale of various processes

Non-linear dependence
of various parameters.

**What is the relevant
timescale when
estimating k?**

**For smaller scales – we
need to include more
processes**





Water-side convection

Turbulent diffusion

$p\text{CO}_{2\text{-atm}}$

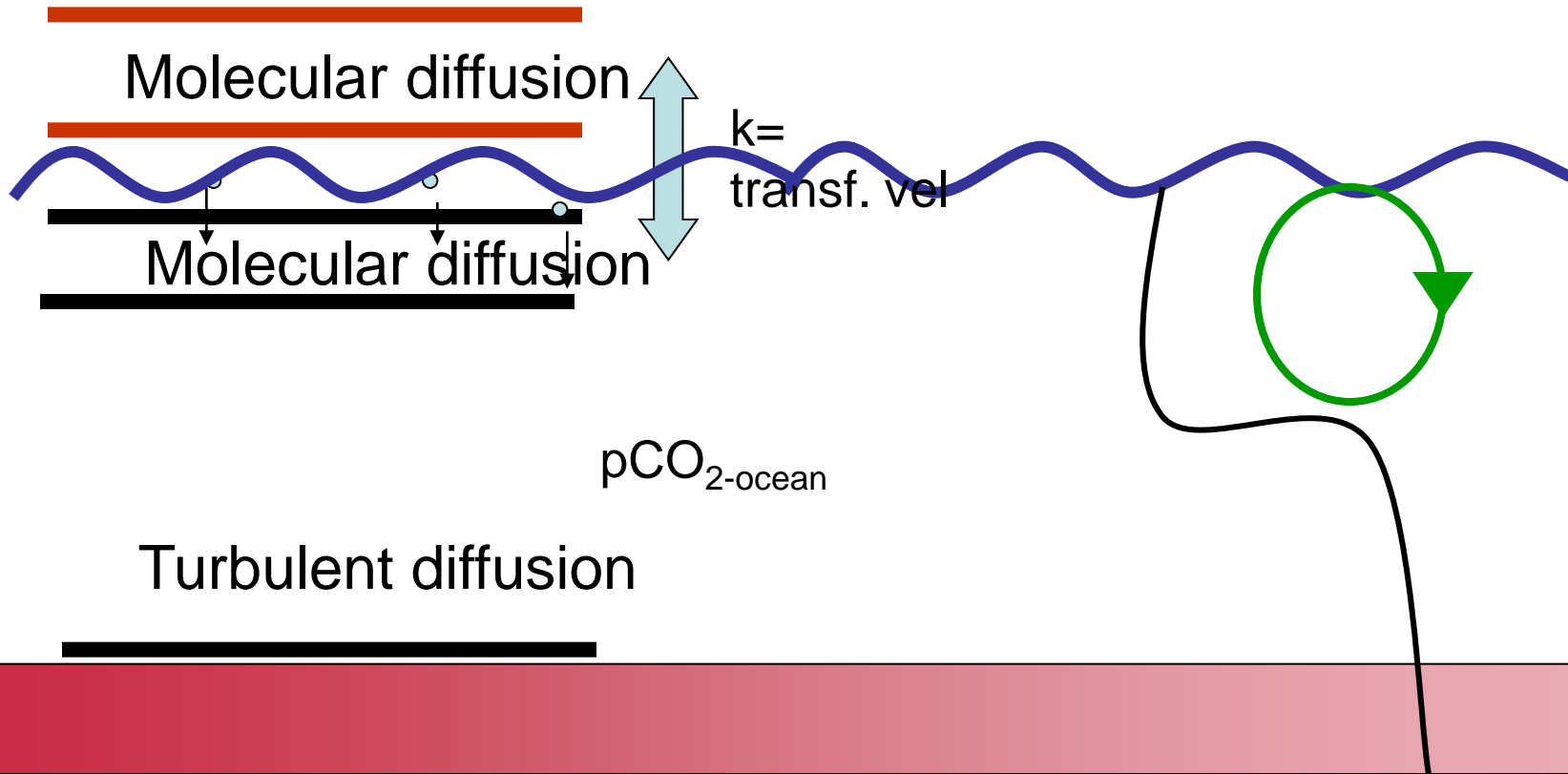
Molecular diffusion

$k =$
transf. vel

Molecular diffusion

$p\text{CO}_{2\text{-ocean}}$

Turbulent diffusion





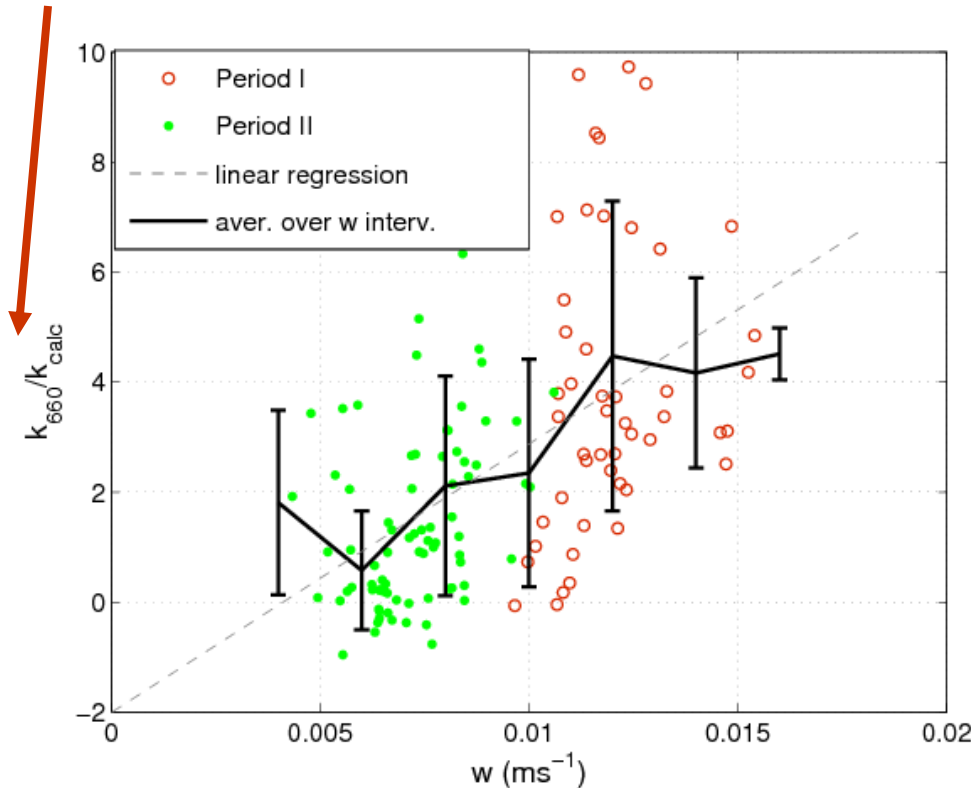
Convective velocity scale

Transfer velocity
normalised
by wind speed

$$w = (z_{ml} B)^{1/3}$$

$B = \text{buoyancy at the surface}$

$z_{ml} = \text{mixed-layer depth}$



B refers to the buoyancy in the water due to cooling and evaporation (colder saltier water is heavier)



How can this be described?

- **Additional resistance in the water** (Rutgersson et al, 2011)

$$\frac{1}{r_w} = \frac{1}{r_{wt} + r_{wm}} + \frac{1}{r_B} + \frac{1}{r_{wc}}$$

$$k = \frac{u_{*a}}{r_w + r_a \alpha}$$

$$\frac{1}{r_{wc}} = \gamma \sqrt{\frac{u_{*w}}{w}} = \textit{enhancement by convection}$$

Strength of buoyancy relative to shear induced turbulence



Methods

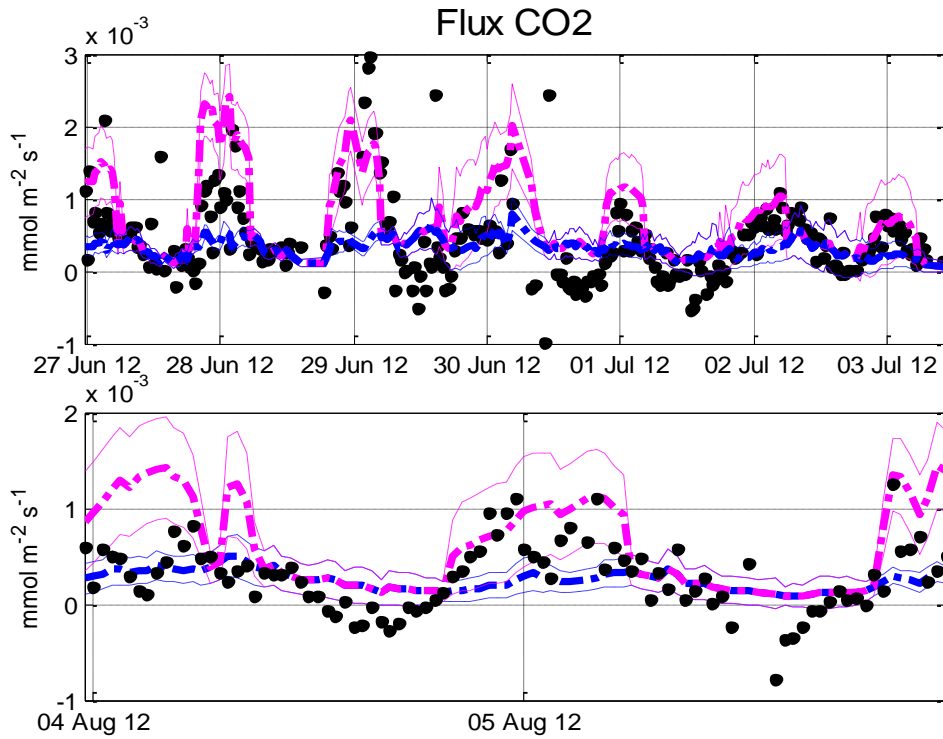
- Different scale basins (Lakes, Baltic Sea)
- Different scalars (carbon dioxide, methane, oxygen, water vapor)



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Lake studies

Strong diurnal cycle for methane and carbon dioxide
(Podgrajsek et al, 2013a; 2013b)



Flux
black=measured
Blue=wind (Cole and Caraco 1998)
Pink=wind+conv
(Rutgersson and Smedman 2010)

Carbon dioxide



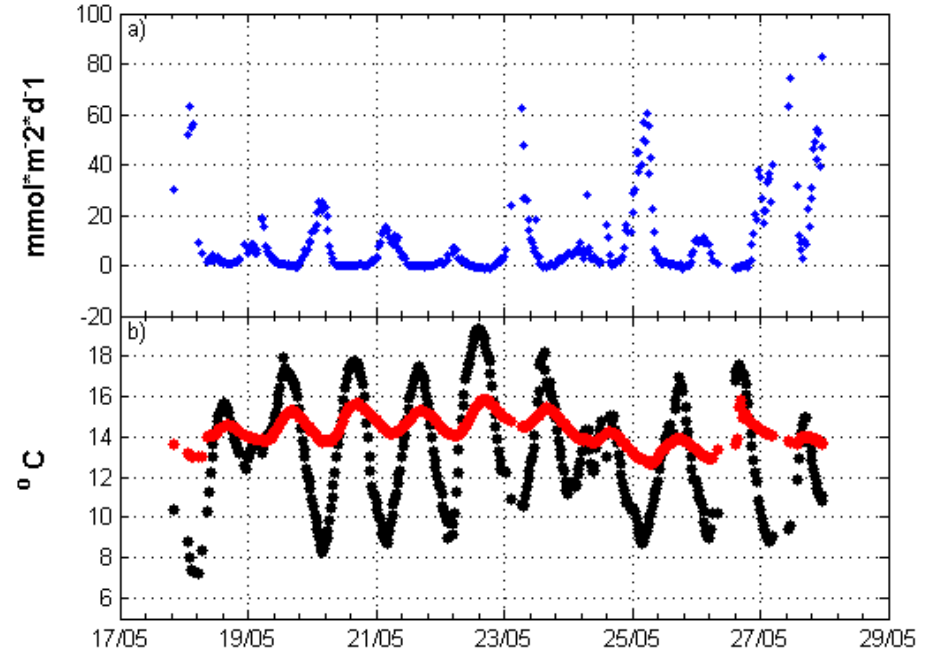
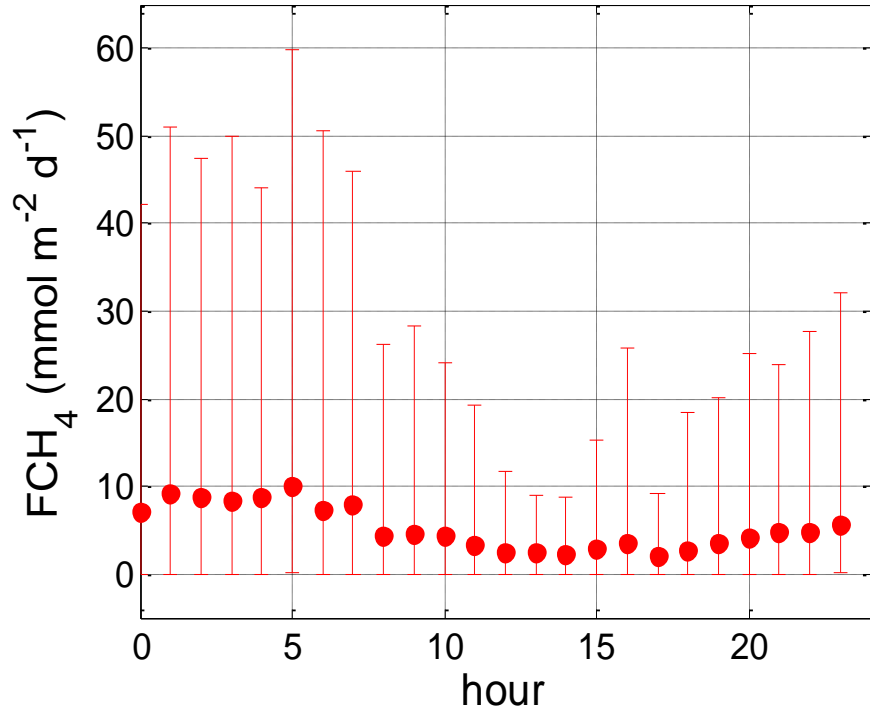


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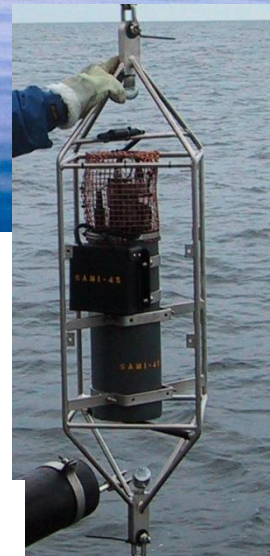
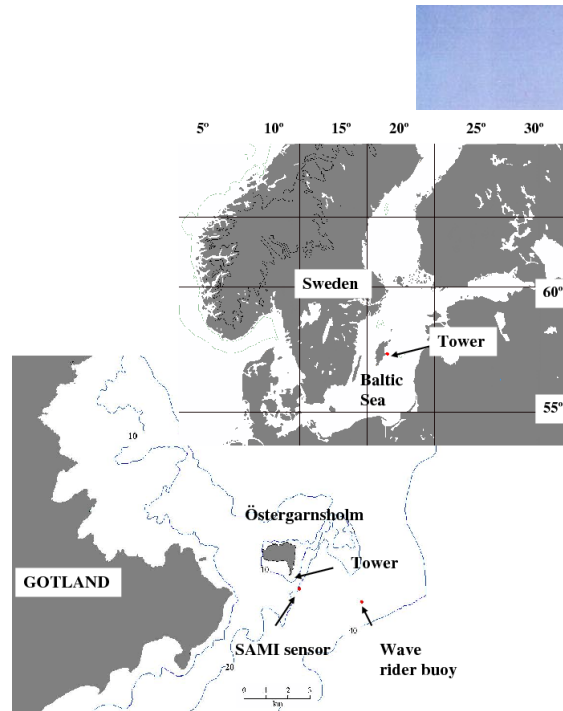
Methane





Baltic Sea studies, the Östergarnsholm-site

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Data for air-sea interaction investigations, meteorological data as well as buoy data (wave, pCO₂ etc). Has been shown to well represent marine conditions.

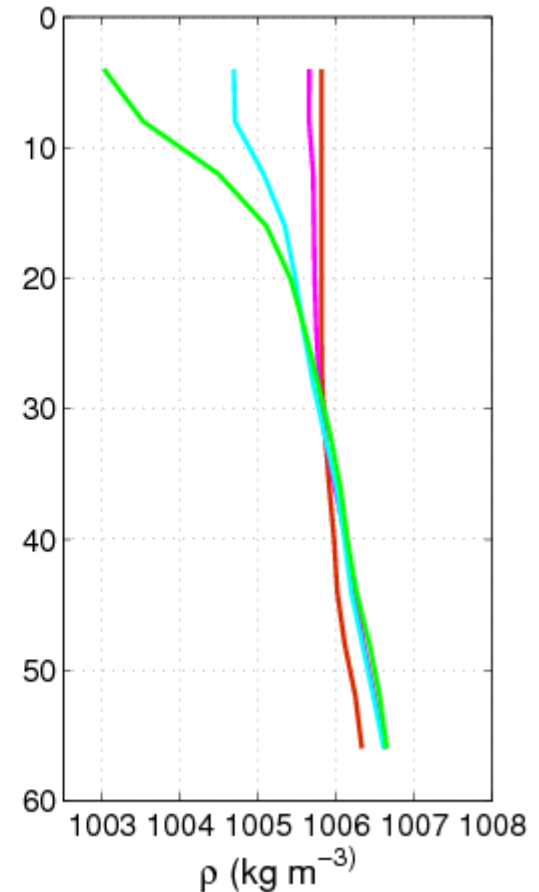
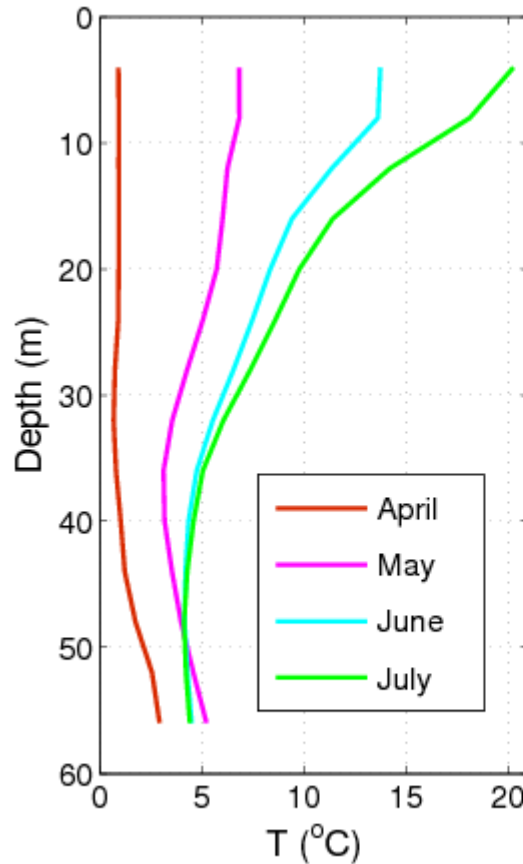


Mixed layer depth

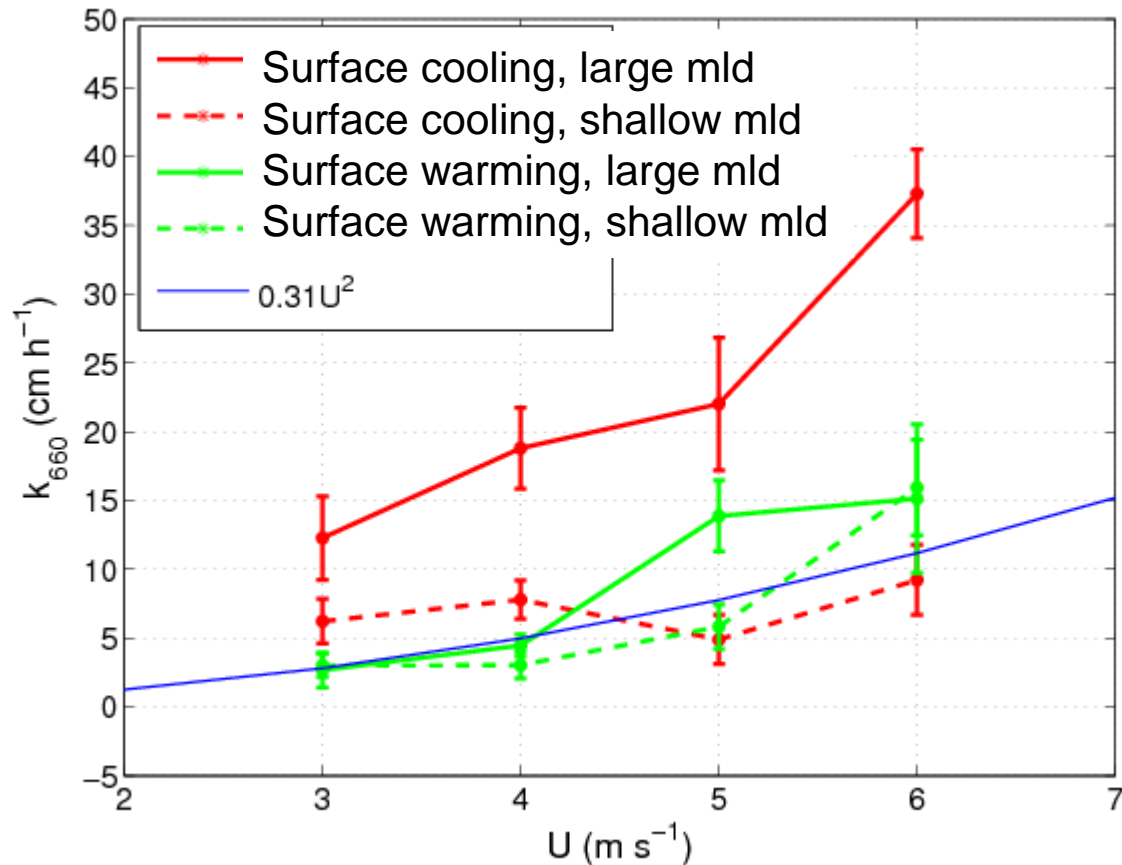
Period I: April to May
($z_{ml} > 40\text{m}$)

Period II: June to July
($z_{ml} < 20\text{m}$)

- Mixed layer depth from HIROMB (3D ocean model for Baltic Sea)



Larger transfer velocity for convection in the water and larger mixed layer depth



Important for intermediate wind speeds

Transfer velocity for varying solubility. EC measurements of Oxygen (Presens MicroTX).

Larger k for higher winds (preliminary results)



Final comments

- When we increase the resolution in time and space – we need knowledge of the relevant processes.
- Impact of surface heat flux underestimated (buoyancy as forcing parameter). Impact also seen in the open ocean (for example GASEX01, McGillis et al, 2004).
- Other gases, additional knowledge



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THANK YOU!

References:

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McGillis, W. R., et al. (2004), Air-sea CO₂ exchange in the equatorial Pacific, *J. Geophys. Res.*, 109, C08S02, doi:10.1029/2003JC002256.

Podgrajsek, E., Rutgersson, A. and E. Sahlée, 2013a. Diurnal cycle of lake methane fluxes. *Geophys. Res. Lett.* Under revision

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